

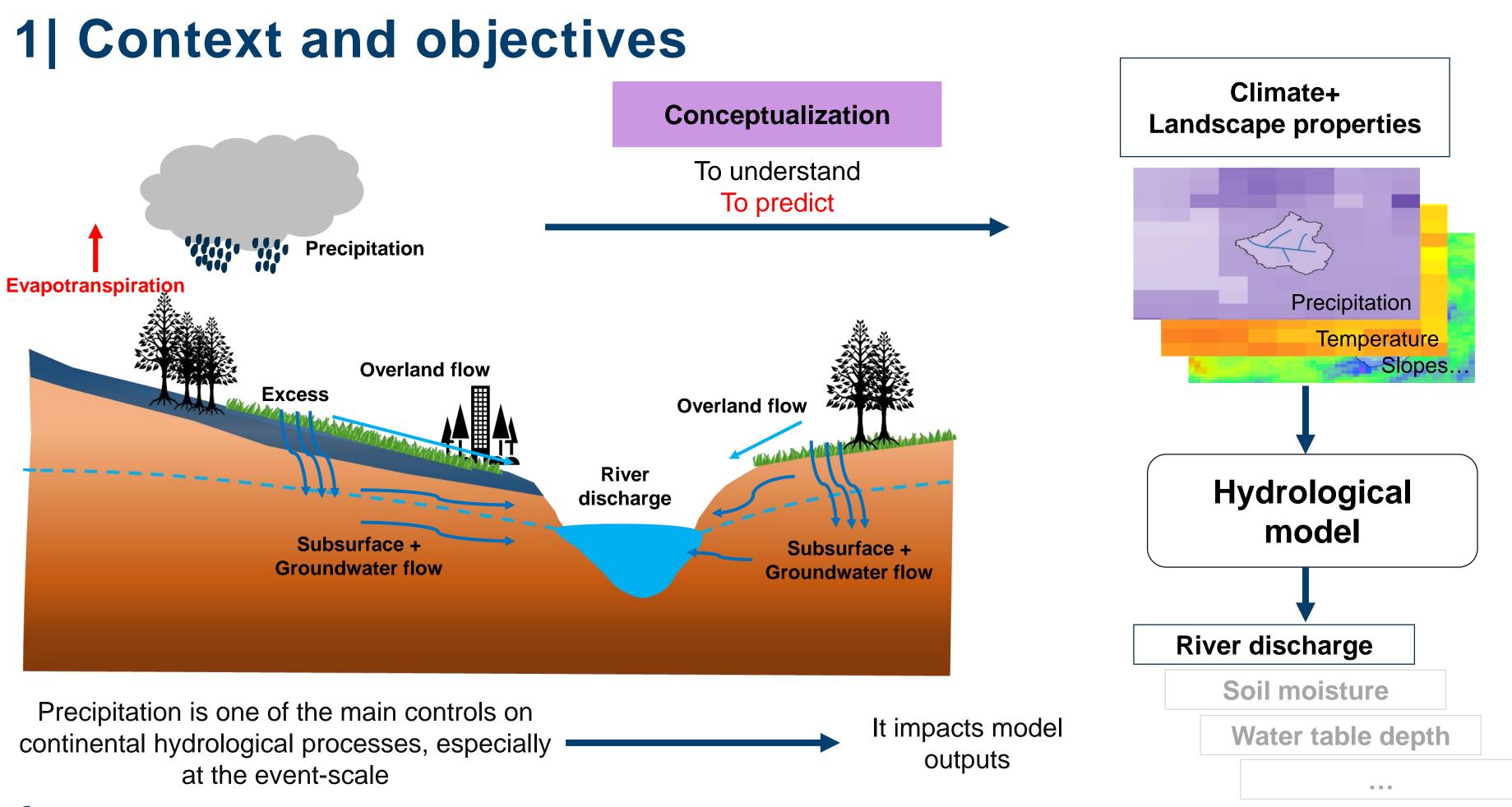
Towards hydrological validation of radar-based precipitation estimates and nowcasts M. Saadi^{1,2}, C. Furusho-Percot^{1,2}, A. Belleflamme^{1,2}, J.-Y. Chen³, R. Reinoso-Rondinel³, S. Trömel^{3,4}, S. Kollet^{1,2} ¹FZJ/IBG-3, ²Geoverbund ABC/J/HPSC-TerrSys, ³UniBonn/Dept. of Meteorology, ⁴Geoverbund ABC/J/CPEX-Lab | P1, P2, P4

2022-04-28 | RealPEP meeting



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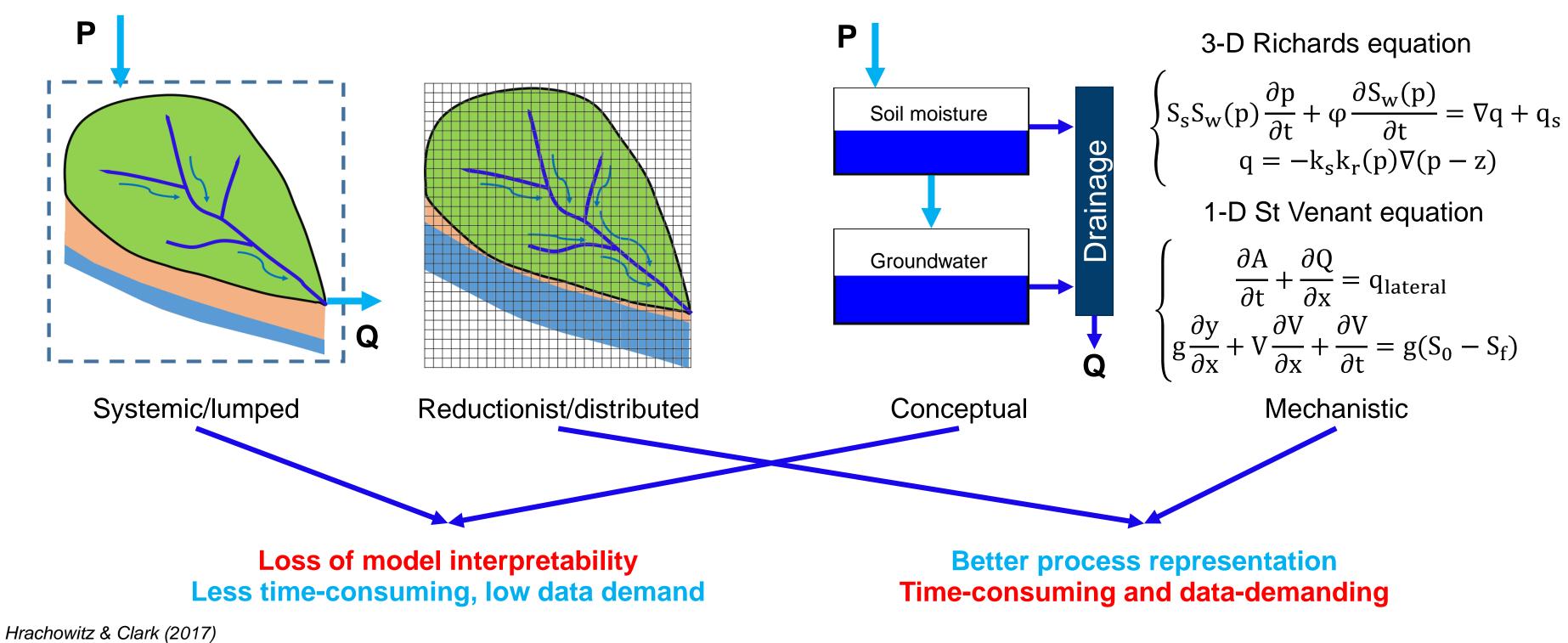






1 Context and objectives

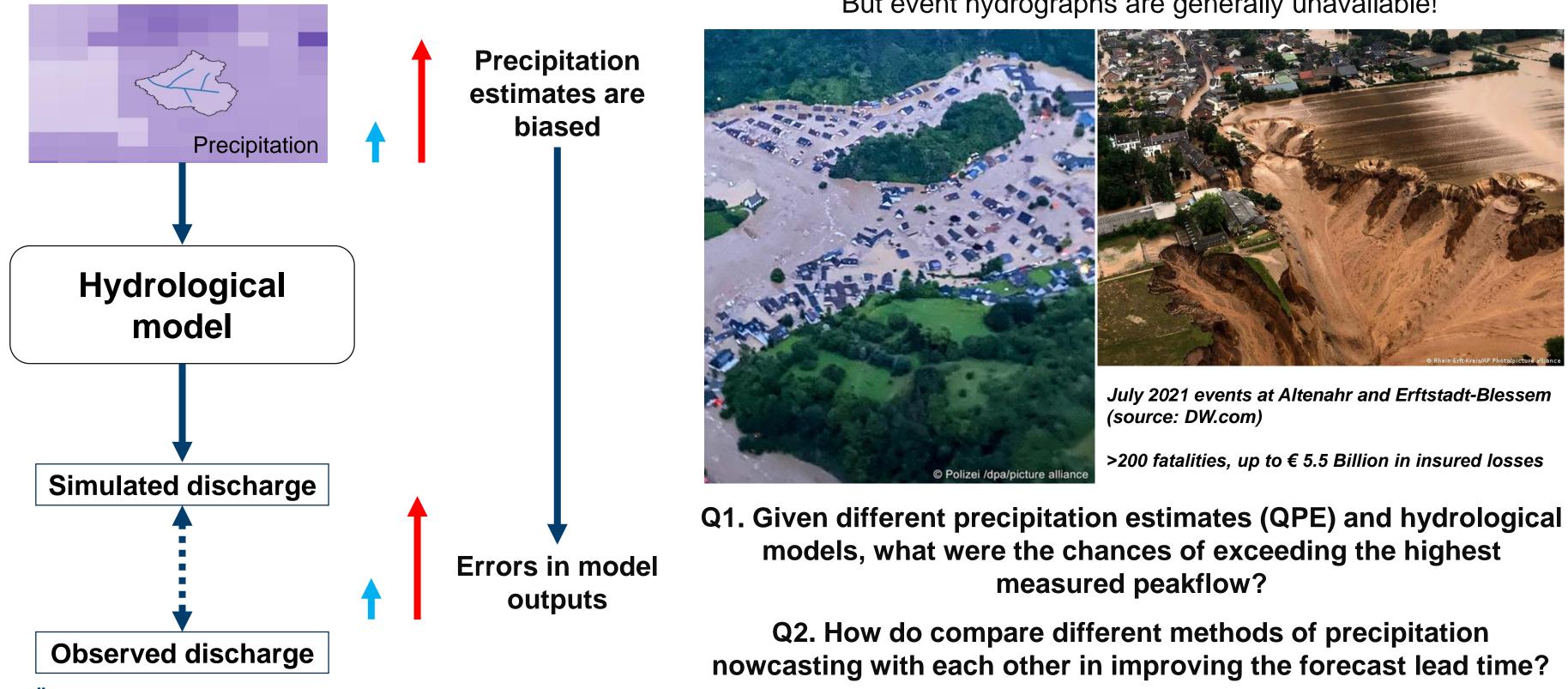
A hydrological model can be:



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1 Context and objectives

Hydrological models can be used to check the accuracy of precipitation estimates

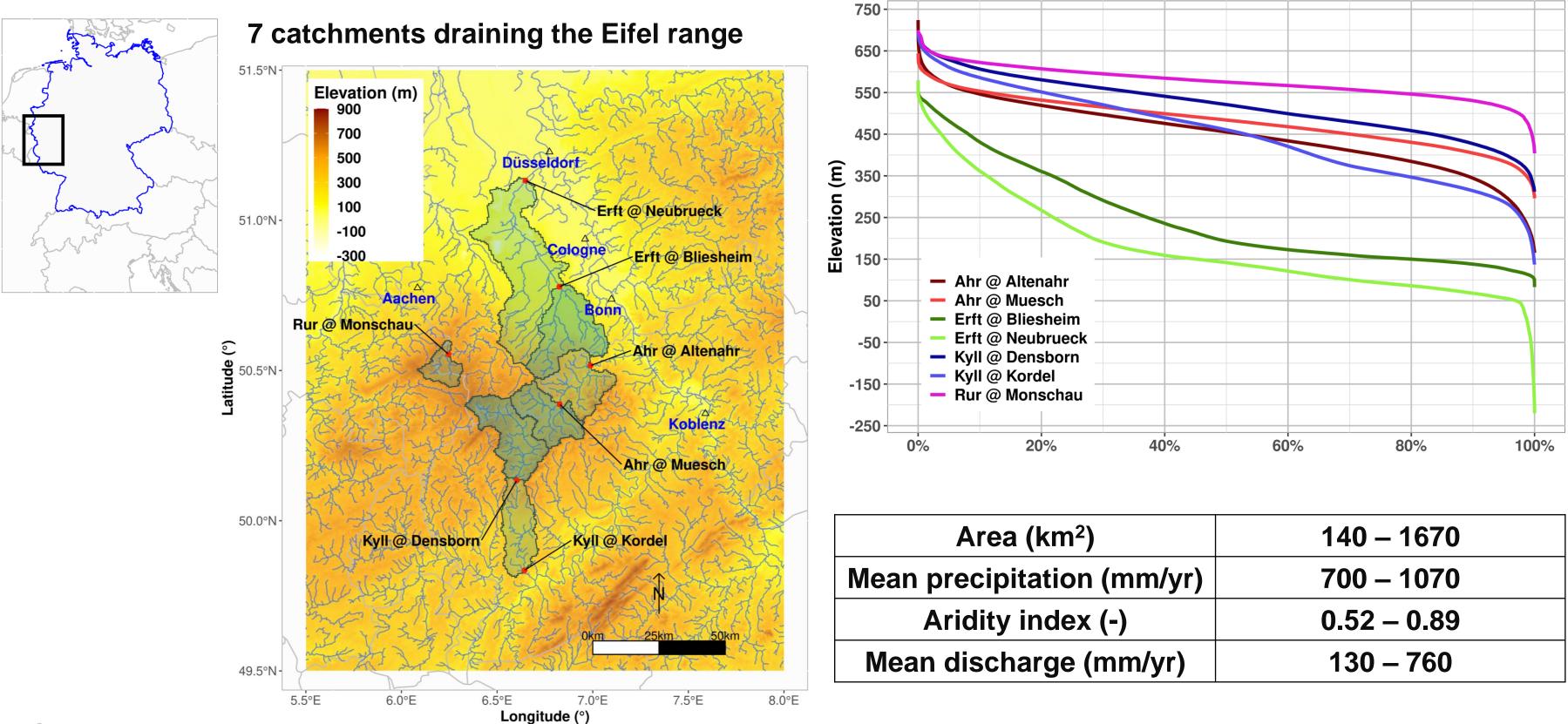


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For extreme floods, accurate precipitation estimates are crucial But event hydrographs are generally unavailable!

2 Catchments, models and data

2.1 | Catchments



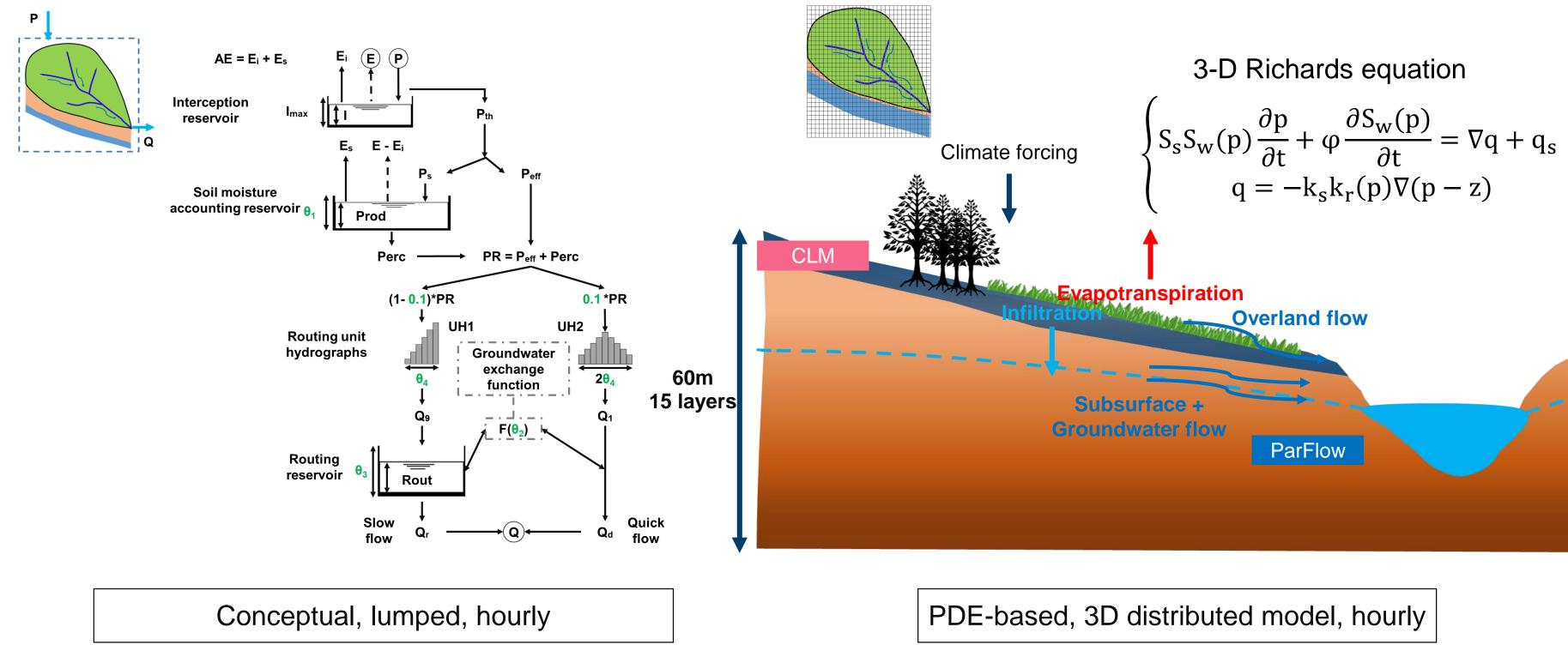
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ı (km²)	140 – 1670
itation (mm/yr)	700 – 1070
index (-)	0.52 – 0.89
arge (mm/yr)	130 – 760

2 Catchments, models and data

2.2 | Models

GR4H (Ficchì et al. 2019)





ParFlow-CLM (Kollet & Maxwell, 2006, 2008)

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2 Catchments, models and data

2.3 | Data

GR4H (Ficchì et al. 2019)

Catchment-averaged inputs

- Precipitation (RADOLAN)
- 2-m air temperature (ERA5-LAND)

ParFlow-CLM (Kollet & Maxwell, 2006)

Cell-averaged inputs (for 2000x2000x15 cells over Central Europe, 611m resolution)

- 2-m air temperature (ERA5-LAND)
- Surface pressure (ERA5-LAND)
- (ERA5-LAND)

Catchment-averaged parameters

- 4 parameters, calibrated using discharge data (LANUV-NRW, LfU-RLP), 2007-2021
- Calibration needs definition of objective function and period of calibration \rightarrow 12 optimal parameter sets for each catchment

Cell-averaged parameters

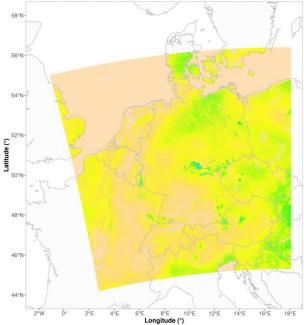
- Topography: ASTER+MERIT DEMs
- and IHME
- 3 tested Manning's n

0.2 s/m^{1/3} (HMann) 0.1 s/m^{1/3} (MMann)

Runs on GPUs of the JUWELS HPC system (4 nodes x 512 GiB)

Runs on local computer

- Precipitation (RADOLAN & ERA5-LAND) - 10-m u and v wind components(ERA5-LAND) - Surface solar/thermal radiation downwards



- Land cover: CLC2018, reclassed in 18 IGBP types - Soil types: SoilGrids250m, grouped into 12 USDA classes

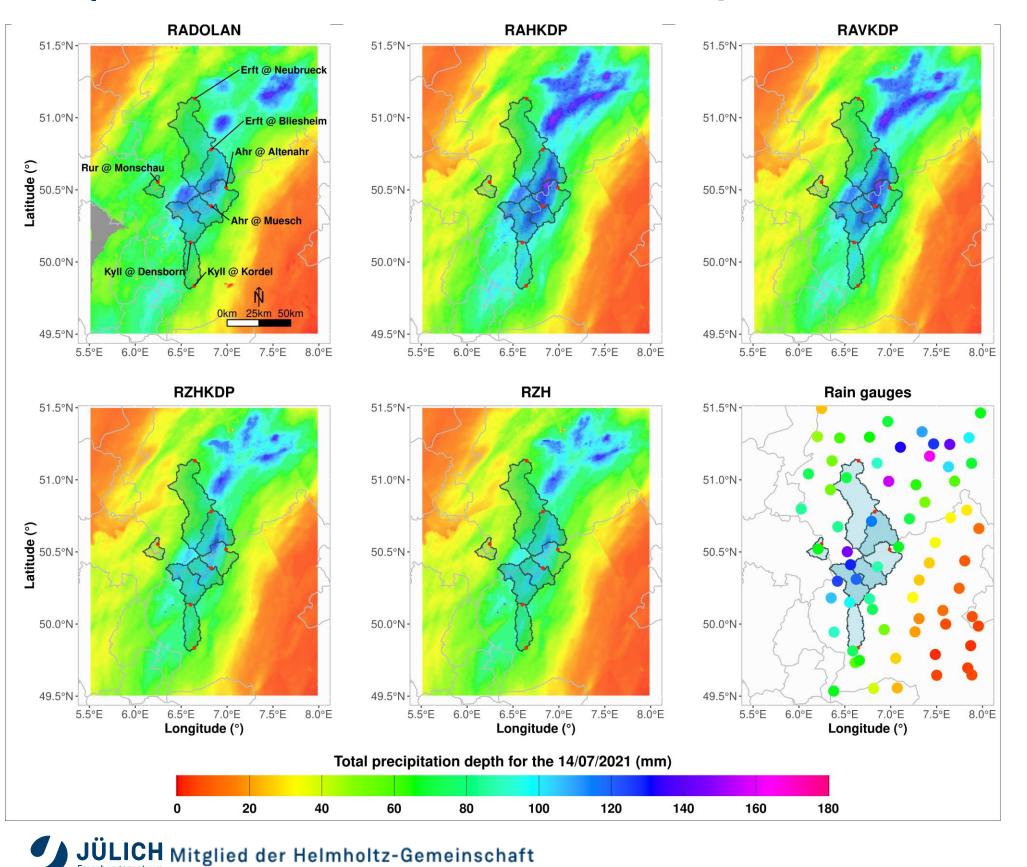
0.02 s/m^{1/3} (LMann)

3 Q1. Impact of QPE & modeling choices on peakflow 3.1 | QPE products for the 14.07.2021 **QPE** products Hydrological models **Peakflow estimations** 12 Q_{psim}(GR4H, Rain gauges) Rain gauges 12 Q_{psim}(GR4H, RADOLAN)+ **GR4H** RADOLAN x12 3 Q_{psim}(ParFlowCLM, RADOLAN) 12 Q_{psim}(GR4H, RZH)+ 3 Q_{psim}(ParFlowCLM, RZH) Horizontal **RZH** reflectivity Z_h 12 Q_{psim}(GR4H, RZHKDP)+ 3 Q_{psim}(ParFlowCLM, RZHKDP) Horizontal reflectivity RZHKDP $Z_{\rm h}$ + Specific differential phase K_{DP} ParFlowCLM **x3** Specific attenuation A_H 12 Q_{psim}(GR4H, RAHKDP)+ RAHKDP + Specific differential 3 Q_{psim}(ParFlowCLM, RAHKDP) phase K_{DP} 12 Q_{psim}(GR4H, RAVKDP)+ 3 Q_{psim}(ParFlowCLM, RAVKDP) Specific attenuation A_v RAVKDP + Specific differential phase K_{DP}



Count the number of times simulations exceeded the historical peakflow for each QPE

3.2 | Result 1: Differences between QPE products

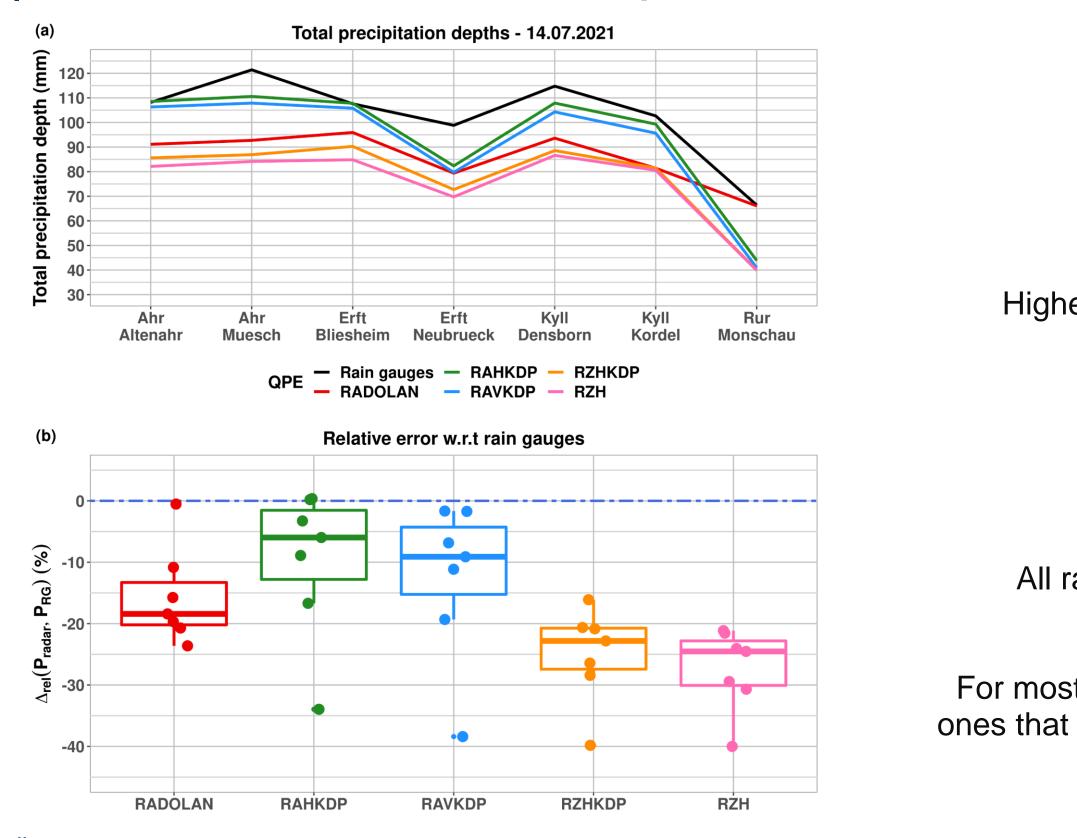


At the pixel scale

Similar spatial pattern

Higher rainfall rates for **RAHKDP** and **RAVKDP**

3.2 | Result 1: Differences between QPE products



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At the pixel scale

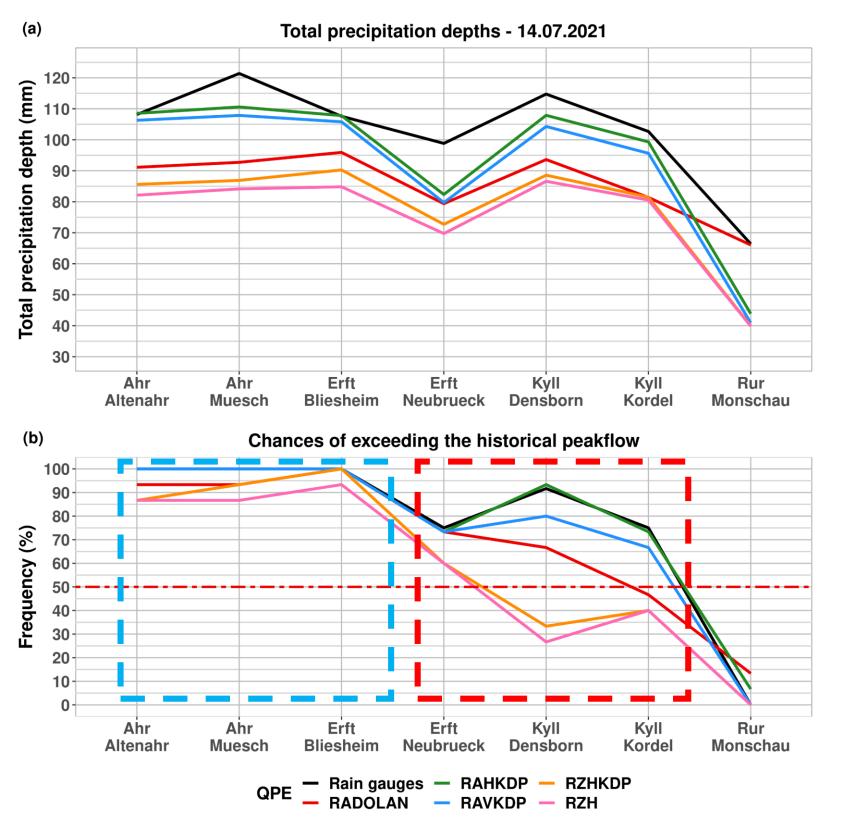
Similar spatial pattern

Higher rainfall rates for **RAHKDP** and **RAVKDP**

At the catchment scale

- All radar-based QPE underestimated the total precipitation from rain gauges
- For most catchments, **RAHKDP** and **RAVKDP** are the ones that agreed most with estimations from rain gauges

3.2 | Result 2: Chances of breaking the historical records of peakflow



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Catch

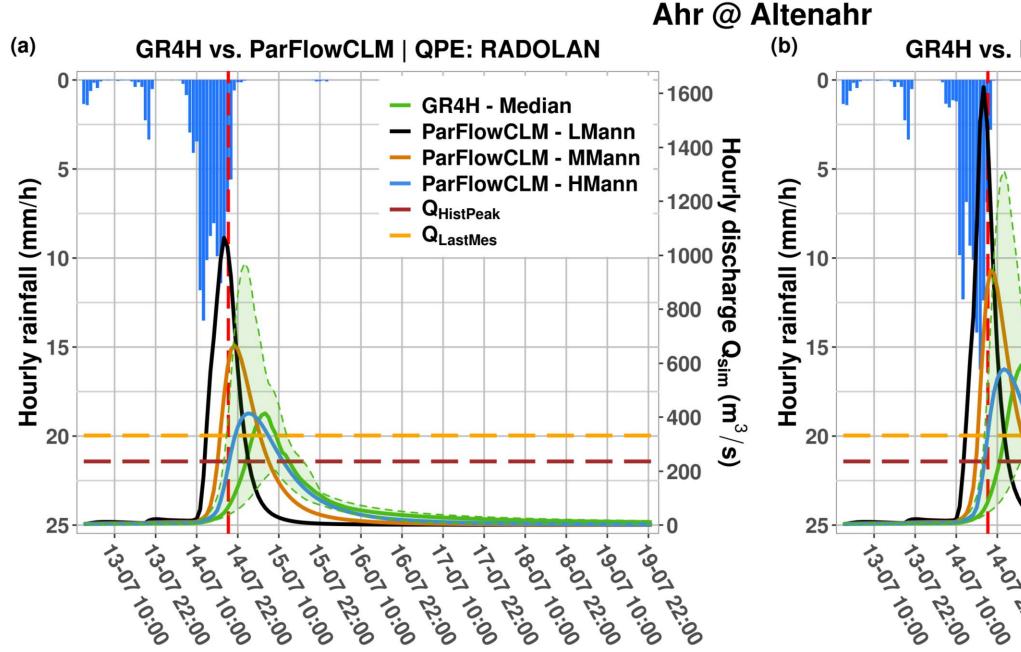
Ahr @ A Ahr @ N Erft @ Bli Erft @ Ne Kyll @ De Kyll @ Rur @ Mo

The effect of the QPE on the (simulated) severity of the event varied among catchments

ment	Historical peakflow (m ³ /s)
ltenahr	236
/luesch	132
iesheim	55.8
eubrueck	46.64
ensborn	180
Kordel	218
onschau	109.63

1. Very high chances no matter what QPE product is used 2. Very low chances for the Rur @ Monschau 3. High dependency on QPE for the remaining catchments

3.2 | Result 3: Uncertainties from model parameters vs. from QPE products



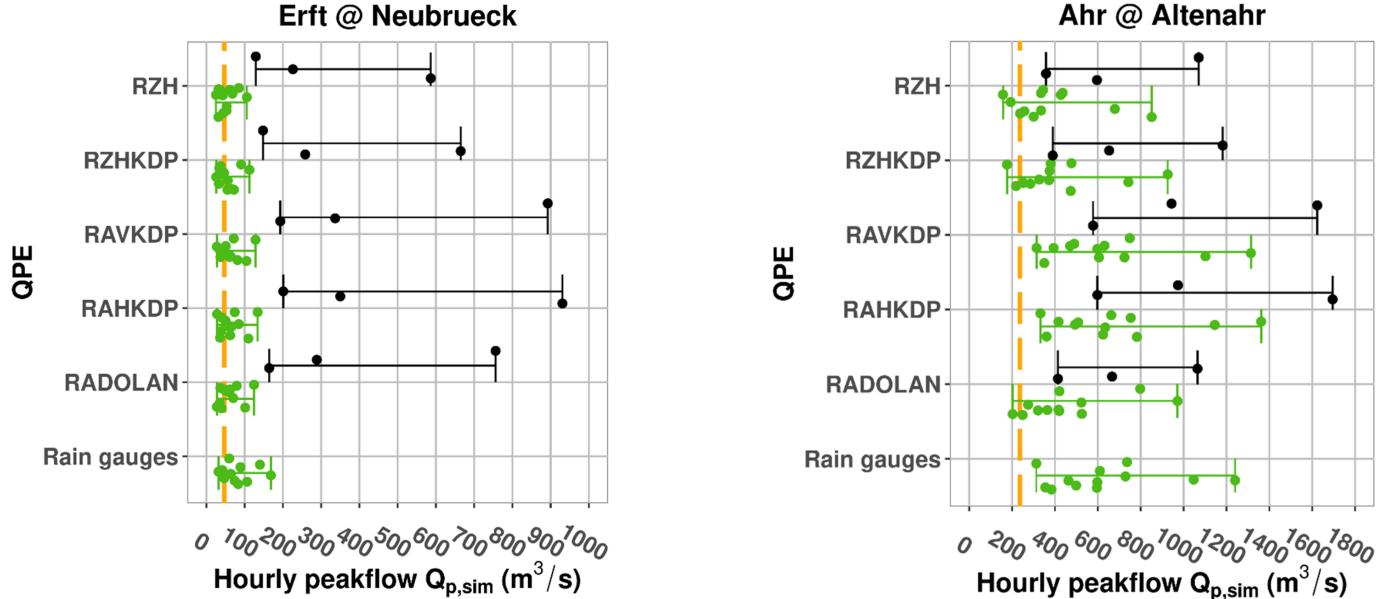
Additional uncertainty from model parameter estimation GR4H simulations bracketed only HMann and MMann



GR4H vs. ParFlowCLM | QPE: RAVKDP 1600 **GR4H** - Median - ParFlowCLM - LMann 1400 Hourly ParFlowCLM - MMann ParFlowCLM - HMann **Q**_{HistPeak} discharge — Q_{LastMes} 1000 -800 **Q**sim -600 (**m**³) -400 S -200 -0 14-07 22:00 16-07 22:00 19-07 10:00 19-07 22:00 15-07 10:00 15-07 22:00 17.07 10:00 17.07 22:00 16-07 10:00 18-07 10:00 18-07 22:00

3 Q1. Impact of QPE & modeling choices on peakflow 3.2 | Result 3: Uncertainties from model parameters vs. from QPE products



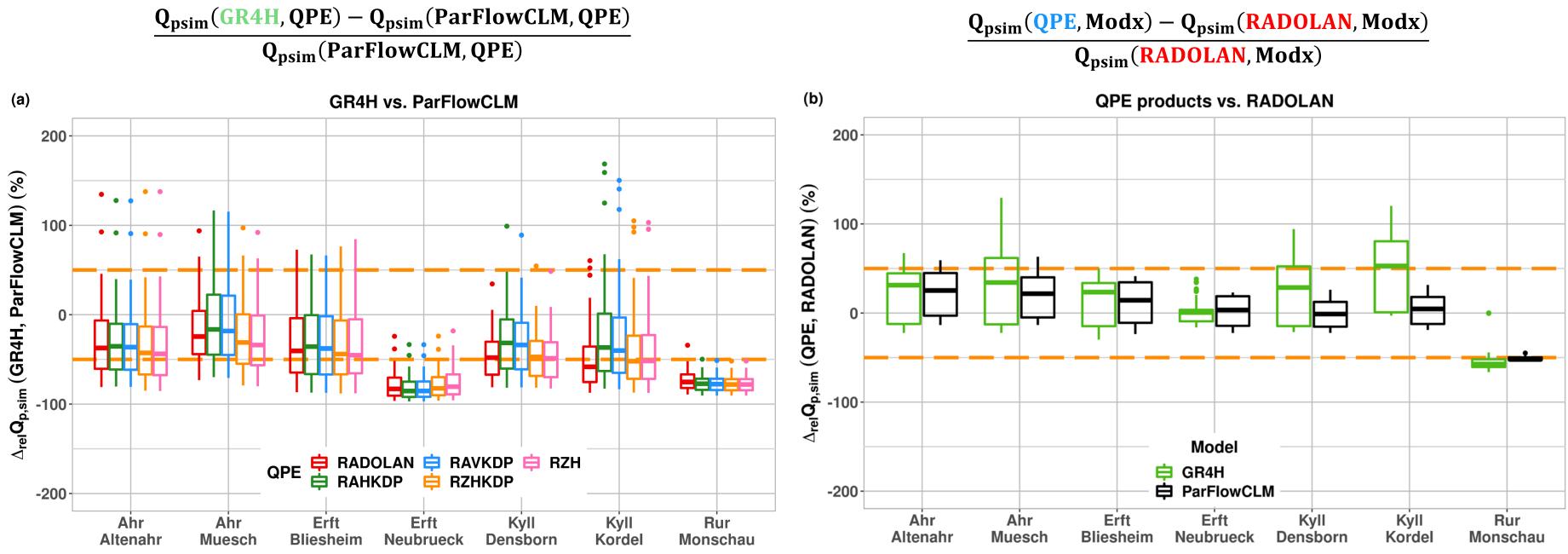


Model disagreement due to anthropogenic influence ?



Slightly influenced catchment

3.2 | Result 3: Uncertainties from model parameters vs. from QPE products



If we change from **ParFlowCLM** to **GR4H**, we tend to have lower peakflow estimates

If we change from **RADOLAN** to any other **QPE**, we tend to have higher peakflow estimates

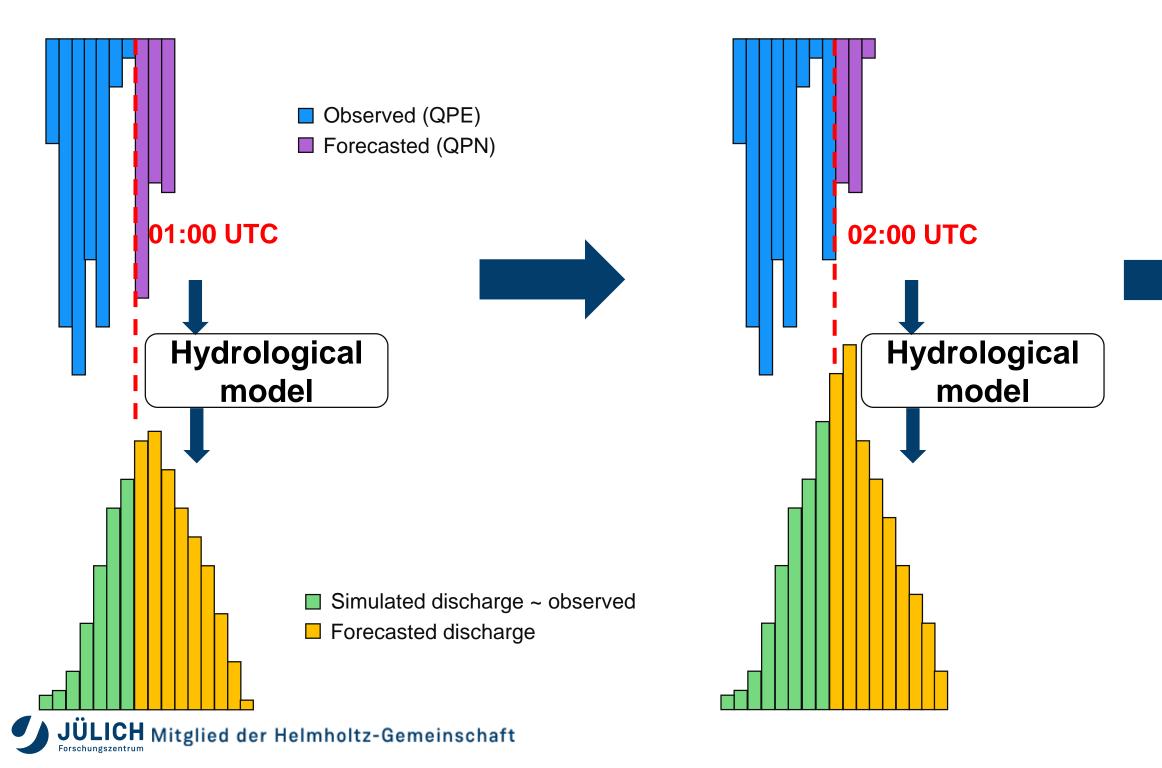


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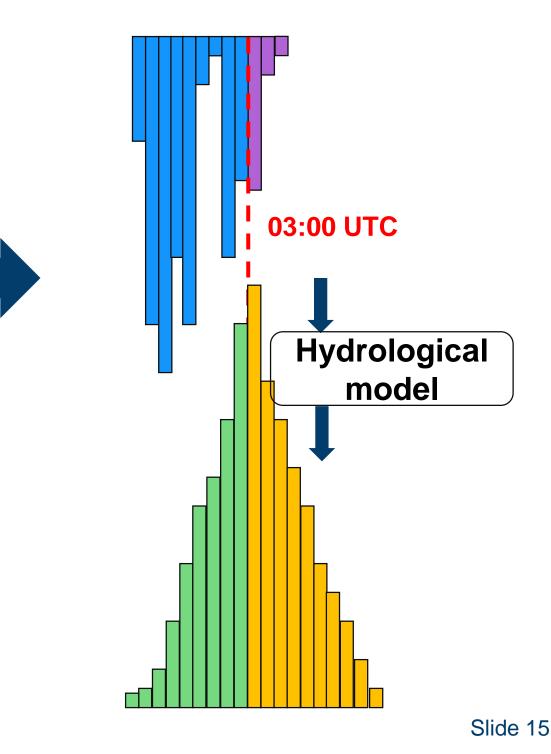
4.1 | QPN methods and framework for hydrological evaluation

Based on the QPE product RAVKDP

2 deterministic: Advection and SPROG (Seed, 2003) + 1 stochastic, with 20 members : STEPS (Bowler et al., 2006)



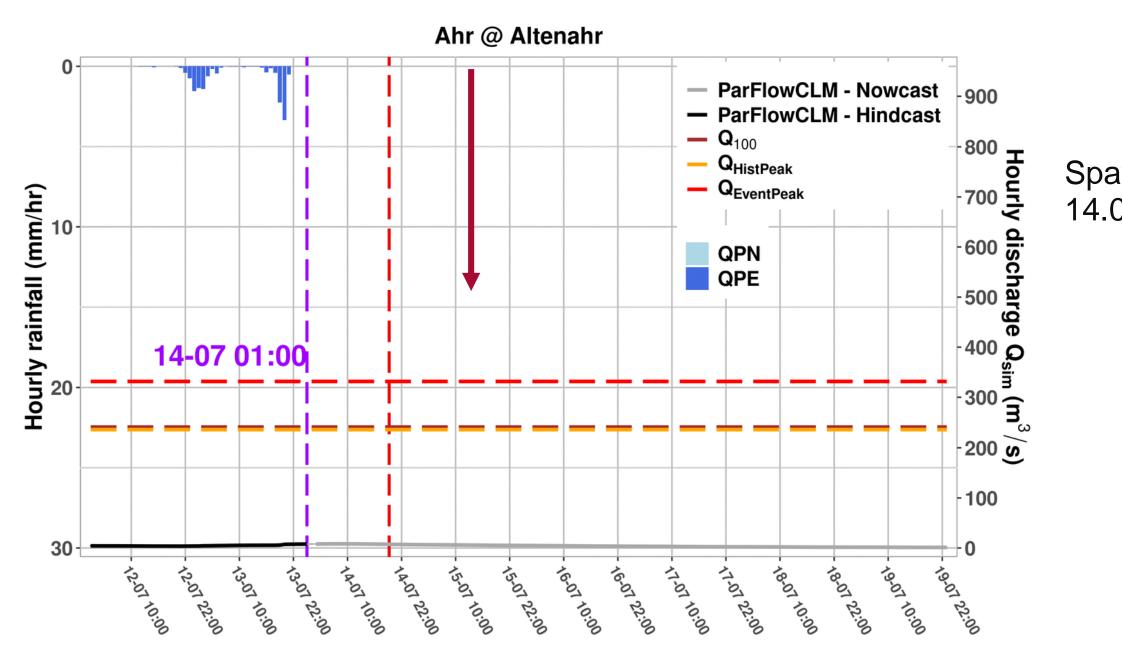




QPN methods and framework for hydrological evaluation 4.1

Based on the QPE product RAVKDP

2 deterministic: Advection and SPROG (Seed, 2003) + 1 stochastic, with 20 members : STEPS (Bowler et al., 2006)



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Spawned each hour between 01h00 and 18h00 of 14.07.2021

QPN methods and framework for hydrological evaluation

Based on the QPE product RAVKDP

2 deterministic: Advection and SPROG (Seed, 2003) + 1 stochastic, with 20 members : STEPS (Bowler et al., 2006)

Evaluation: construct a virtual forecasted hydrograph for each lead time

Time	Q, LT = 1h		Q, LT = 2h	Q,
01h00	Q(1h00) spawned	l at 00h00	Q(1h00) spawned at 23h00 j-1	Q
02h00	Q(2h00) spawned	l at 01h00	Q(2h00) spawned at 00h00	Q
03h00	Q(3h00) spawned	l at 02h00	Q(3h00) spawned at 01h00	Q
04h00	Q(4h00) spawned	l at 03h00	Q(4h00) spawned at 02h00	Q
05h00	Q(5h00) spawned	l at 04h00	Q(5h00) spawned at 03h00	Q
06h00	Q(6h00) spawned	l at 05h00	Q(6h00) spawned at 04h00	Q

Virtual because assembled out of different hydrographs



1, LT = 4h

(1h00) spawned at 21h00 j-1

(2h00) spawned at 22h00 j-1

(3h00) spawned at 23h00 j-1

(4h00) spawned at 00h00

(5h00) spawned at 01h00

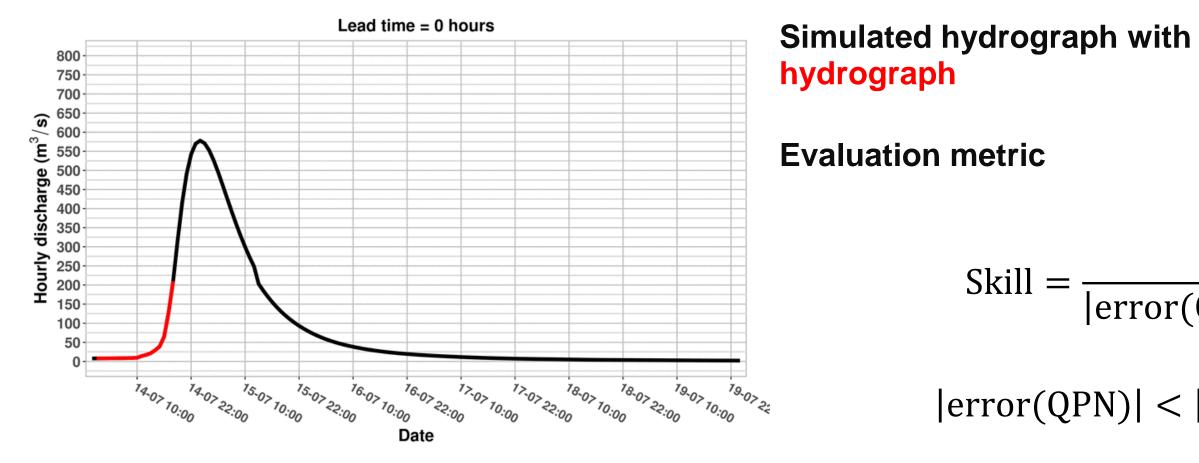
(6h00) spawned at 02h00

QPN methods and framework for hydrological evaluation

Based on the QPE product RAVKDP

2 deterministic: Advection and SPROG (Seed, 2003) + 1 stochastic, with 20 members : STEPS (Bowler et al., 2006)

Evaluation: construct a virtual forecasted hydrograph for each lead time



Chen et al. (2017); Hersbach (2000)

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- Simulated hydrograph with QPE to be compared with forecasted

- $Skill = \frac{|error(Bench)|}{|error(QPN)| + |error(Bench)|}$
- $|error(QPN)| < |error(Bench)| \Rightarrow Skill > 0.5$
- Error estimated using the continuous rank probability score (CRPS)

4.2 | Result 1: Hydrological persistence as benchmark

No differences between the three methods

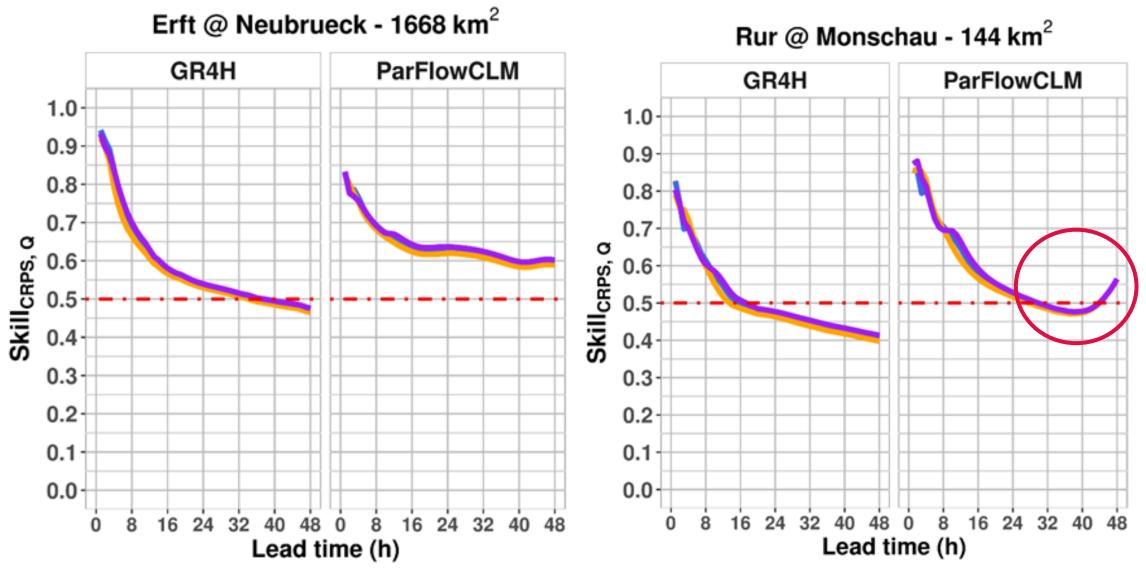
Decreasing skill with increasing lead time

A rebound in the skill curves for the Ahr river. related to rising/falling limbs of the hydrograph

Usefulness that can last as long as 48 h, except for some cases

No significant relation with catchment size

No significant effect of the hydrological model



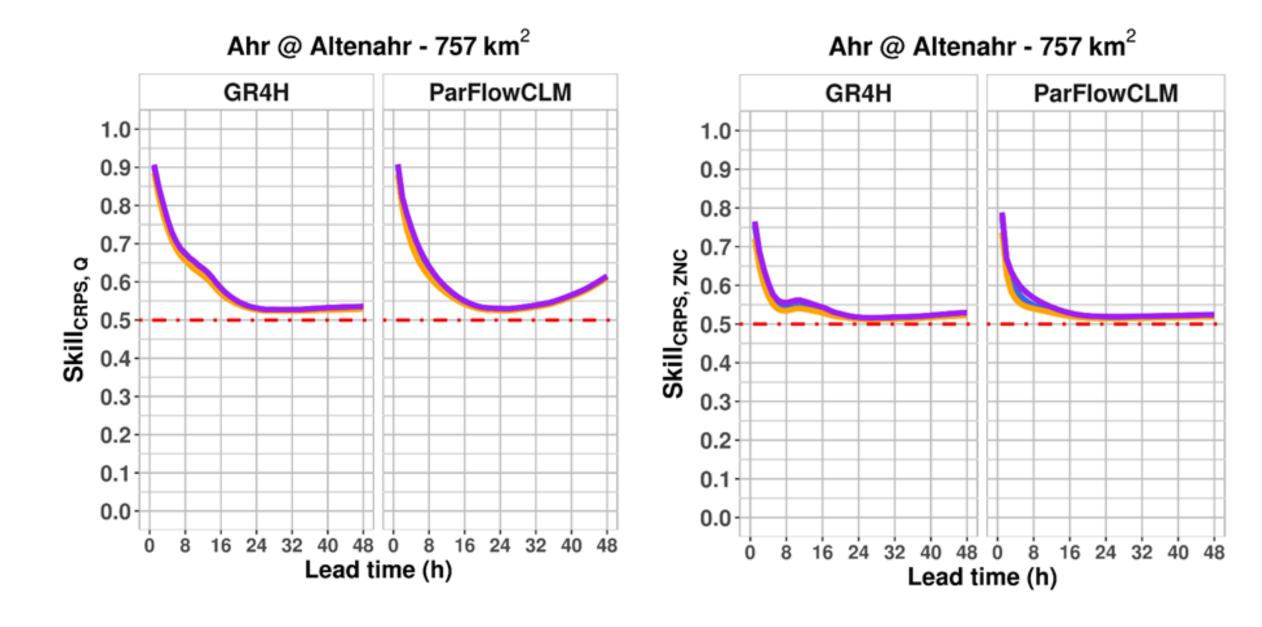


QPN method – Advection – SPROG – STEPS

4.2 | Result 2: Zero-precipitation nowcasts as a benchmark

Similar conclusions, but no rebound effect!!

Lower skill compared to hydrological persistence, but more costly with additional model run



QPN method - Advection - SPROG - STEPS



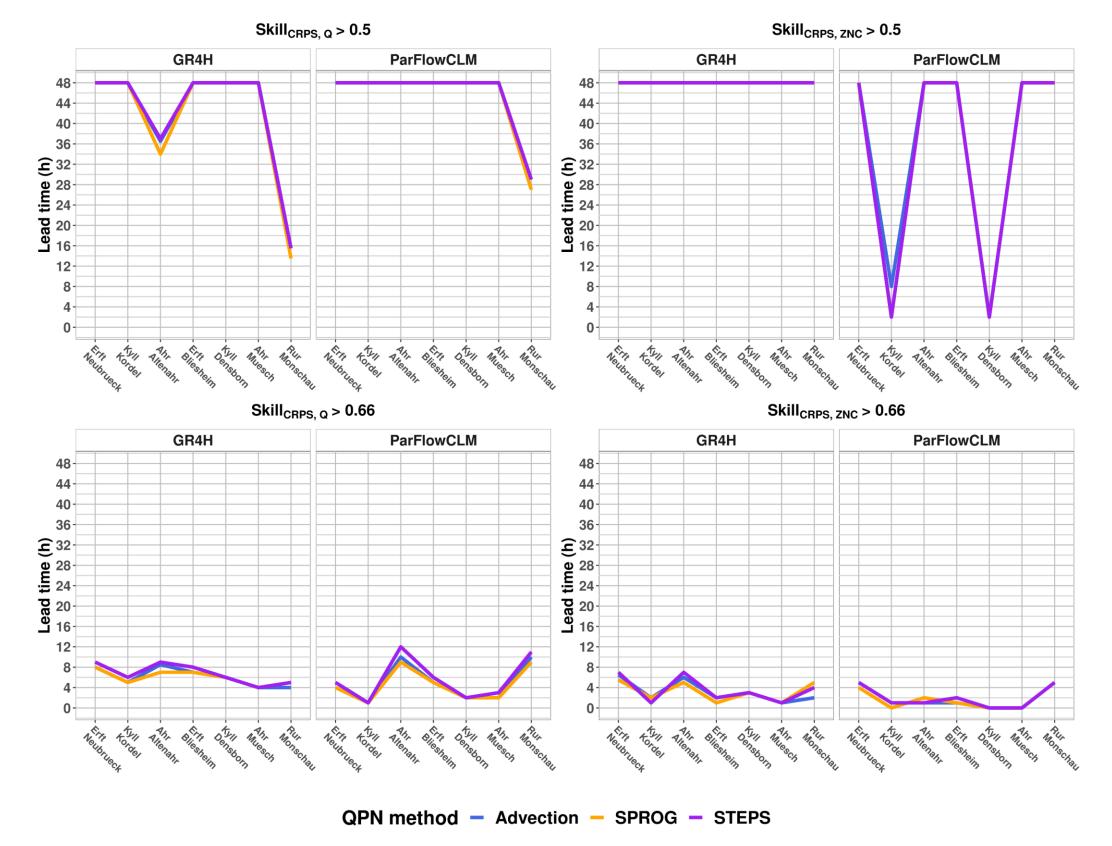
4.2 | Result 3: Effect of skill threshold

 $|error(QPN)| < |error(Bench)| \Rightarrow Skill > 0.5$

$$|\operatorname{error}(\operatorname{QPN})| < \frac{1}{2} |\operatorname{error}(\operatorname{Bench})| \Rightarrow \operatorname{Skill} > \frac{2}{3}$$

Changing the skill threshold has a significant impact!

QPN are now useful only up to 12 h at most!



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5 Conclusions

Comparison of modeling philosophies

There is general agreement between GR4H and ParFlowCLM, except for highly influenced catchments. Adopting a very low Manning's explains most of the discrepancies between the models

At this stage, running a conceptual model seems more advantageous, but the distributed model allows for estimations in all points of the domain, regardless of discharge data availability

Evaluation of QPE products

Including specific attenuation helped improve the radar-based QPE products

The choice of QPE products impacted the ability of models to anticipate a record-breaking flood

Added value of QPN methods

The QPN methods behaved similarly. Possible differences in a less predictable, more spatially variable precipitation event? The choice of the benchmark model and the skill threshold impacts the evaluation of the QPN



Thank you for your attention! **Questions?**

References

Saadi, M., Furusho-Percot, C., Belleflamme, A., Chen, J.-Y., Trömel, S., and Kollet, S., in review, 2022. How uncertain are precipitation and peakflow estimates for the July 2021 flooding event? Nat. Hazards Earth Syst. Sci. Discuss. [preprint], https://doi.org/10.5194/nhess-2022-111

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