

PARA - **PArameterization informed by RAdar** - **Update** T. Scharbach^{1,2}, S. Trömel¹

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photo of Bonn X-band radar by V. Pejcic

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Filtering method example



Figure 1: Horizontal reflectivity Z_H (left) monitored with BoXPol on 30 May 2016 at 03:06 UTC together with the profile of calculated minimum normalised Shannon information entropy (right). The applied threshold is indicated as red line and the excluded sequences displayed as transparent regions.

- Filtering approach applied to synthetic and observation data:
 - ightarrow 1168.5 hours of BoXPol observations from 627 different days (2014-2018)
 - ightarrow 109.2 hours from 11 different days (2017-2018) for JuXPol + BoXPol synthetic data
- To increase statistical significance, ICON simulations of IWC, D_m and N_t of JuXPol + BoXPol in retrieval space are merged with all simulated C-band stations from DWD

$$D_m(K_{DP}, Z_{dp}, Z_h) = \begin{cases} D_m(K_{DP}, Z_{dp}) & \text{if } Z_{DR} \ge 0.4 \text{ dB} \\ D_m(K_{DP}, Z_h) & \text{if } Z_{DR} < 0.4 \text{ dB} \end{cases}$$
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$$IWC(K_{DP}, Z_{dr}, Z_h) = \begin{cases} IWC(K_{DP}, Z_{dr}) & \text{if } Z_{DR} \ge 0.4 \text{ dB} \\ IWC(K_{DP}, Z_h) & \text{if } Z_{DR} < 0.4 \text{ dB} \end{cases}$$
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$$N_t = N_t(IWC(K_{DP}, Z_{dr}, Z_h))$$
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following [Bukovčić et al., 2020], [Bukovčić et al., 2018], [Ryzhkov and Zrnic, 2019] and [Carlin et al., 2021]

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CFTD of retrieved D_m and simulated D_m



Figure 2: CFTDs of retrieved D_m (left) and simulated D_m (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

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CFTD of retrieved N_t and simulated N_t



Figure 3: CFTDs of retrieved N_t (left) and simulated N_t (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

CFTD of retrieved IWC and simulated IWC



Figure 4: CFTDs of retrieved IWC (left) and simulated IWC (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

CFTD of observed Z_H and synthetic Z_H



Figure 5: CFTDs of observed Z_H (left) and synthetic Z_H (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

CFTD of observed Z_{DR} and synthetic Z_{DR}



Figure 6: CFTDs of observed Z_{DR} (left) and synthetic Z_{DR} (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

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CFTD of observed K_{DP} and synthetic K_{DP}



Figure 7: CFTDs of observed K_{DP} (left) and synthetic K_{DP} (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020].

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CFTD of observed ρ_{HV} and synthetic ρ_{HV}



Figure 8: CFTDs of observed ρ_{HV} (left) and synthetic ρ_{HV} (right), with The solid red line representing the mean and the dashed red lines the 20th and 80th percentiles. The blue line shows the number of samples in a 1 °C layer. Temperature information is taken from ERA5 [Hersbach et al., 2020] and ICON.

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References i

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 $\label{eq:result} \mbox{RADAR}\mbox{MIE}\mbox{LM}\mbox{and}\mbox{RADAR}\mbox{MIE}\mbox{LM}\mbox{and}\mbox{reflectivity}\mbox{ from model output.}$

Technical Report 28, Consortium for Small Scale Modeling (COSMO).

- Bukovčić, P., Ryzhkov, A., and Zrnić, D. (2020).
 Polarimetric relations for snow estimation—radar verification.
 Journal of Applied Meteorology and Climatology, 59(5):991–1009.

Bukovčić, P., Ryzhkov, A., Zrnić, D., and Zhang, G. (2018). **Polarimetric radar relations for quantification of snow based on disdrometer data.** *Journal of Applied Meteorology and Climatology*, 57(1):103–120.

References ii

Carlin, J. T., Reeves, H. D., and Ryzhkov, A. V. (2021).
 Polarimetric observations and simulations of sublimating snow: Implications for nowcasting.

Journal of Applied Meteorology and Climatology, 60(8):1035–1054.

 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., et al. (2020).
 The era5 global reanalysis.

Quarterly Journal of the Royal Meteorological Society, 146(730):1999–2049.

Ryzhkov, A. V. and Zrnic, D. S. (2019).

Radar polarimetry for weather observations.

Springer.

References iii

- Shannon, C. E. (1948).
 A mathematical theory of communication. 27(3):379–423.
- Trömel, S., Ryzhkov, A. V., Hickman, B., Mühlbauer, K., and Simmer, C. (2019).
 Polarimetric radar variables in the layers of melting and dendritic growth at x band—implications for a nowcasting strategy in stratiform rain.
 Journal of Applied Meteorology and Climatology, 58(11):2497–2522.
- Wolfensberger, D., Scipion, D., and Berne, A. (2016).
 Detection and characterization of the melting layer based on polarimetric radar scans.

Quarterly Journal of the Royal Meteorological Society, 142:108–124.