

PRISTINE

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Polarimetric Radar simulations with realistic Ice and Snow properties and multi-frequency consistency Evaluation

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- Motivation

- Making EMVORADO fit for external scattering data
 - Bulk scattering lookup tables from external data sources
 - Orientation averaging

- Outlook

→ Approach: add polarimetry to EMVORADO, but **keep existing core features & characteristics**

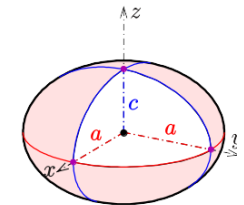
- consistent model coupling, sensor (network) modelling
- hydrometeor property assumptions (e.g. hydrometeor morphology & melting state)
- speed (→bulk scattering lookup-tables)

*state-of-the-art,
but **has its issues***

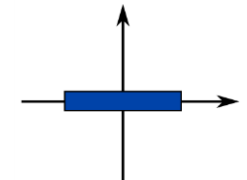
→ Added scattering model option: **T-Matrix + angular moments**

- **default:** shape (AR), orientation (σ_β), melt fraction dependence from **Ryzhkov et al. (2011)**

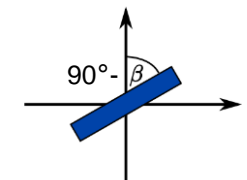
liquid	rain	ice	snow	graupel, hail	
Rayleigh	oblate spheroids	oblate spheroids	oblate spheroids	oblate spheroids	shape
-	Brandes (2002) f(deg4-in-D)	Matrosov (1996) thick plates aD^b	1.0-0.02*D 0.8 (D>10mm)	1.0-0.02*D 0.8 (D>10mm)	AR
-	10°	10°	40°	40°	σ_β
-	-	both: lin. in f_m to rain	both: lin in f_m to rain	AR: lin. in f_m between AR _{wet} =[AR _{rain} , 0.8, 0.48, AR _{rain}] for f_m =[0, 0.8, 1] σ : lin. in f_m to rain	melting behaviour (f_m =mass melt fraction)



Basic orientation

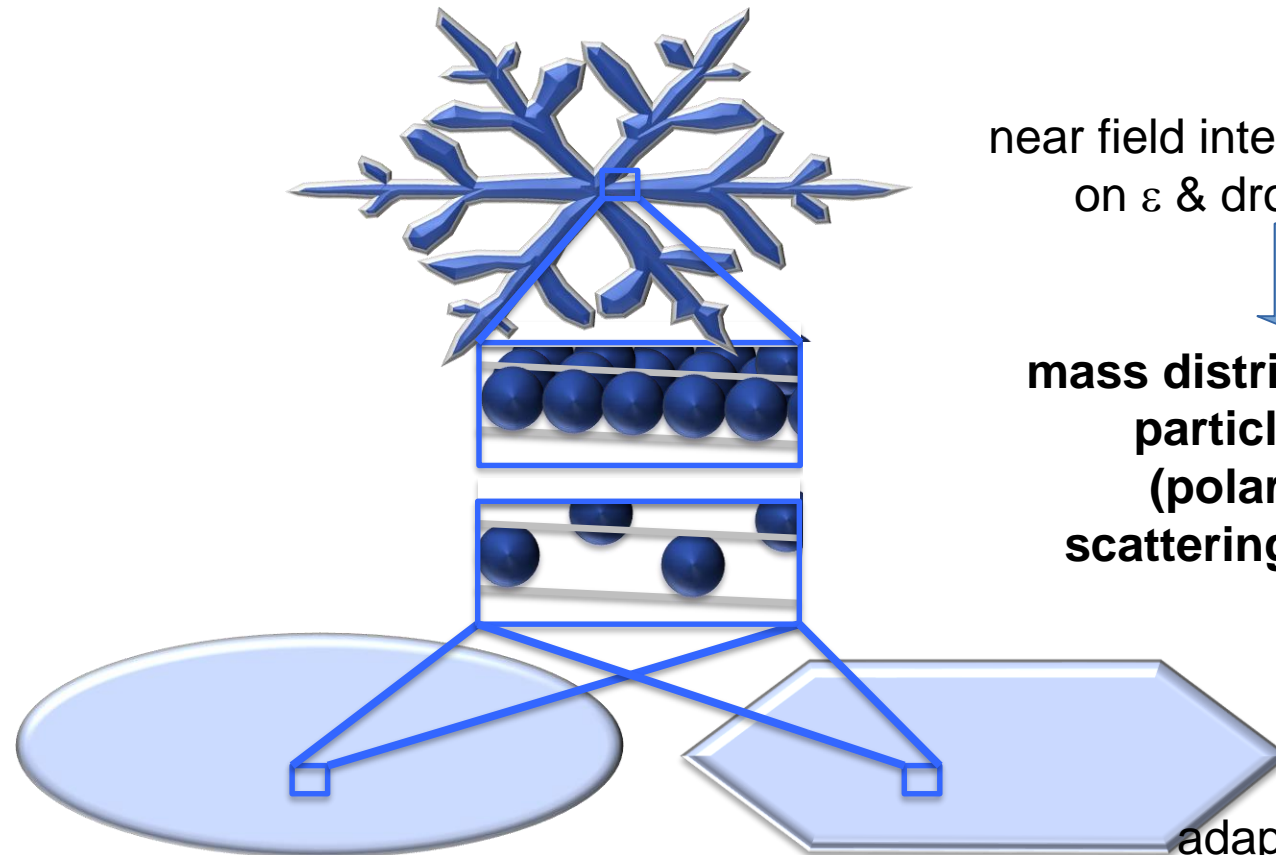
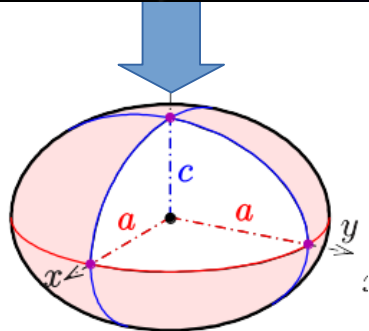
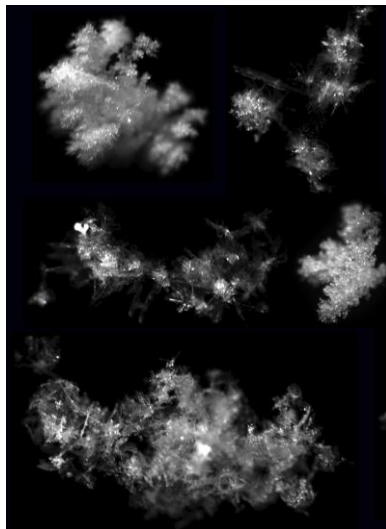


Tilted orientation



→ **Polarimetric** radar forward operator state-of-the-art:

- Assume regular-shaped, **homogeneous effective density** in Mie/**T-Matrix** calculations
- (most?) popular: fix D_{\max} , m , aspect ratio → derive effective (**reduced**) density of spheroid



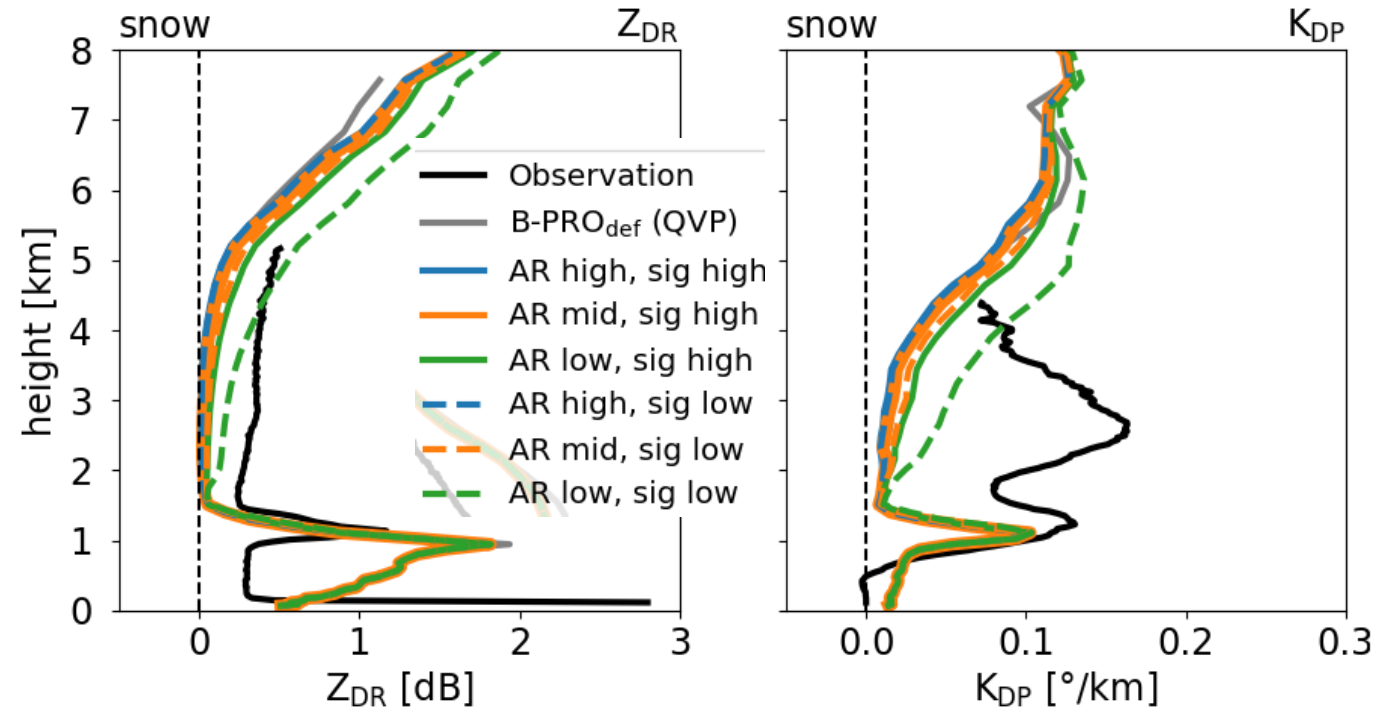
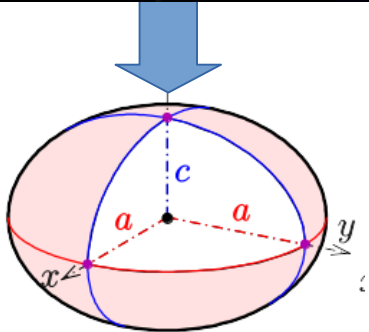
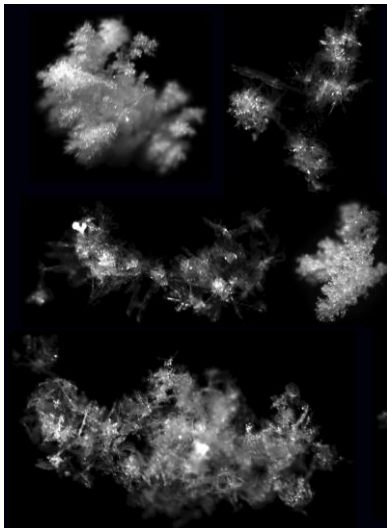
near field interactions depend on ϵ & drop off as $1/r^3$



mass distribution within particle affects (polarimetric) scattering properties

adapted from Kumjian, 2019

- **Polarimetric** radar forward operator state-of-the-art:
- Assume regular-shaped, **homogeneous effective density** in Mie/**T-Matrix** calculations
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Shrestha et al., 2022

- Approach: add polarimetry to EMVORADO, but **keep existing core features & characteristics**
 - consistent model coupling, sensor (network) modelling
 - hydrometeor property assumptions (e.g. hydrometeor morphology & melting state)
 - speed (→bulk scattering lookup-tables)
- Added scattering model option: T-Matrix + angular moments

- Add option to **use externally calculated scattering data**
 - interface to scattering DB(s)

→ Improve DWD's radar operator EMVORADO

- for (better) use in
 - **model evaluation: O-B deviations due to NWP model or radar operator?**
 - data assimilation: bias reduction (at the source, not post-proc)
- but keeping its capabilities
 - model consistency (esp. PSD, m-D relation)
 - calculation speed (eg. bulk scatt. lookup tables)
 - flexibility in instrument definition, e.g. frequency

→ Motivation

→ Making EMVORADO fit for external scattering data

- **Bulk scattering lookup tables from external data sources**
- Orientation averaging

→ Outlook

- Ensure consistency
 - w/ model (PSD, m-D relation)
 - w/ EMVORADO setup (e.g. melting scheme)
- } ensured by file naming scheme using hash-based IDs
no user interaction required (nor possible)
- Phase 1: Externally prepared model-consistent LUT
 - a) **Overrule/don't apply hash ID**
 - b) Simplified hash number: only consider PSD & m-D
 - Phase 2ff: in-EMVORADO LUT preparation
 - Select particles from DB with consistent m-D
 - (& other constraints – here's the science part)
 - Form & use equivalent hash ID

- Motivation

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- Outlook

- So far: **Angular moments approximation** (Ryzhkov, e.g. 2011)
- Derive polarimetric scattering properties of particles with Gaussian canting distribution from scattering amplitudes of „uncanted“ particle
 - **Pro:**
 - single orientation output from T-Matrix only (minimum effort & memory)
 - „quasi-analytical“ orientation weight factors (easy computation)
 - **Con:**
 - rotationally symmetric particles only (?)
 - easy-to-calculate factors for elevation=0° and specific canting distributions only (oblate: random & Gaussian; prolate: horizontally aligned)

→ Modification: ***Explicit orientation averaging***

- Derive polarimetric scattering properties for a set of fixed particle orientation & integrate/average over them
- **Con:**
 - more T-Matrix output (more memory; **comp. effort increases only marginally when done correctly**)
 - numerical integration, ie method and sampling point dependent
- **Pro:**
 - any desired orientation distribution (flexible)
 - any radar elevation (flexible)
 - arbitrary particle shapes (flexible)

→ replace angular moments, also for in-EMVORADO SSP calculations from T-Matrix

→ Parameters to provide & average:

- Angular moments act on complex scattering amplitudes **S**
 - **S** are NOT additive! (ie not usable in expl. orient. avg.!)
- currently, set of additive parameters used in PSD-integration, hydrometeor class add-up, and bulk LUT:
 - **zh** (=bsc-xs in h-h), **zv**, **zvh**, **rrhv** (=Re(nominator(RHV))), **irhv** (=Im(...)); **kdp**, **ah** (=ext-xs in h), **adp**

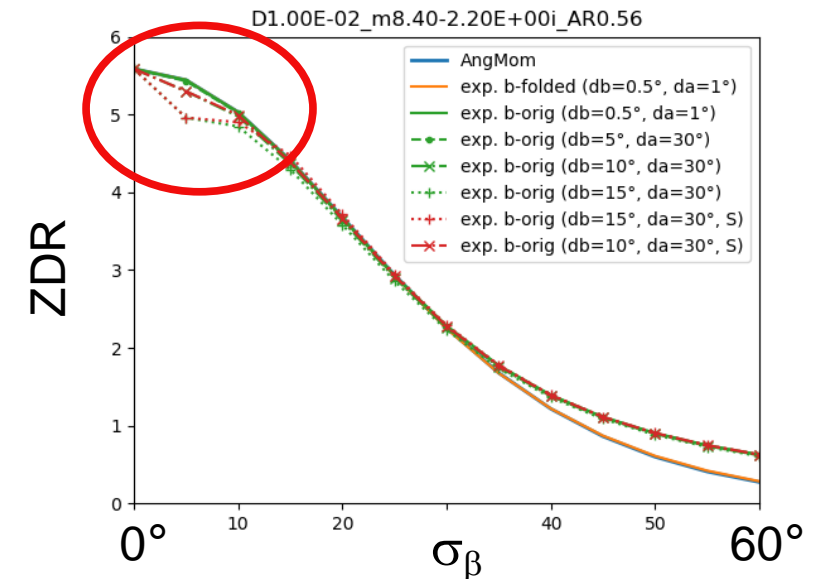
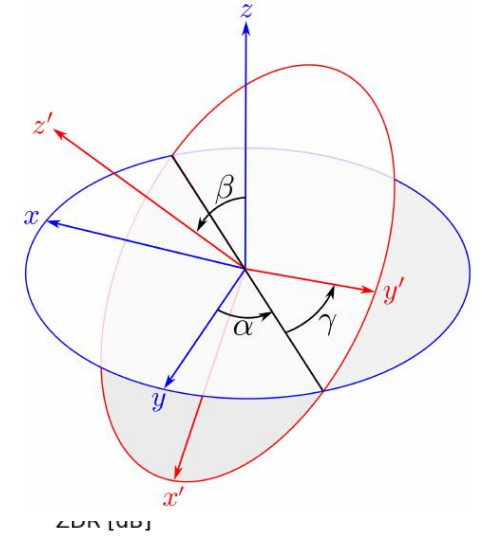
$$S_{hv} = \frac{Z^2 |S_{hh} S_{vv}^*|}{\sqrt{|S_{hh}|^2 \cdot |S_{vv}|^2}} = \frac{\sqrt{(\bar{Z}_{33} + \bar{Z}_{44})^2 + (\bar{Z}_{43} - \bar{Z}_{34})^2}}{\sqrt{\sigma_{hh} \cdot \sigma_{vv}}}$$

- (Stokes) Scattering and extinction matrices (Z and K) are an alternative representation of pol. scatt props
 - additive
 - bsc=f(Z(1:2,1:2)), nom(RHV)=f(Z(3:4,3:4)) & ext,kdp=f(K)

→ use LUT 8-parameter set in canting distrib calc & (already) as new TMat-(wrapper-)output

→ Numerical integration:

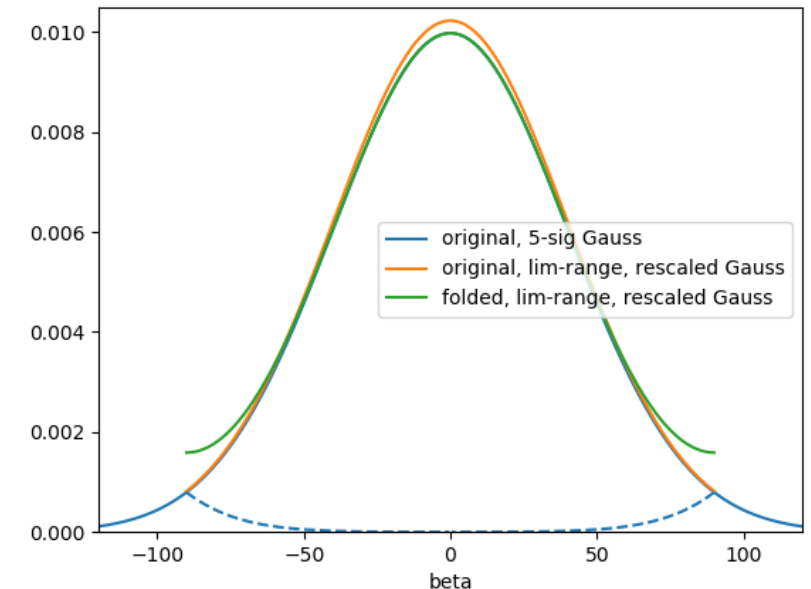
- Over azimuthal (α) and canting (β) orientation angles (for oblate spheroids, γ not applicable due to rotat. symmetry w.r.t. z-axis)
- Which quadrature scheme?
 - Trapezoidal, Simpson; Gauss-Legendre, ...?
- Which polar angle parameter to integrate over?
 - β or $\cos(\beta)=\mu$?
 - $\int Z * p * \sin\beta d\beta = \int Z * p d\mu$
- How many quadrature points?
 - not very critical, but β (variable) more than alpha (α)



→ Numerical integration: Over **canting** (β) orientation angles

- **What about pdf-contributions from outside $\pm 90^\circ$?**
 - for large(r) σ_β , $\text{norm}(\text{Gaussian pdf})$ over $[-90^\circ, +90^\circ] < 1$.
 - What to do?
 - **nothing?** (not an option since not norm-conserving)
 - **renormalize?** (current implementation)
 - **fold?** (my earlier choice)

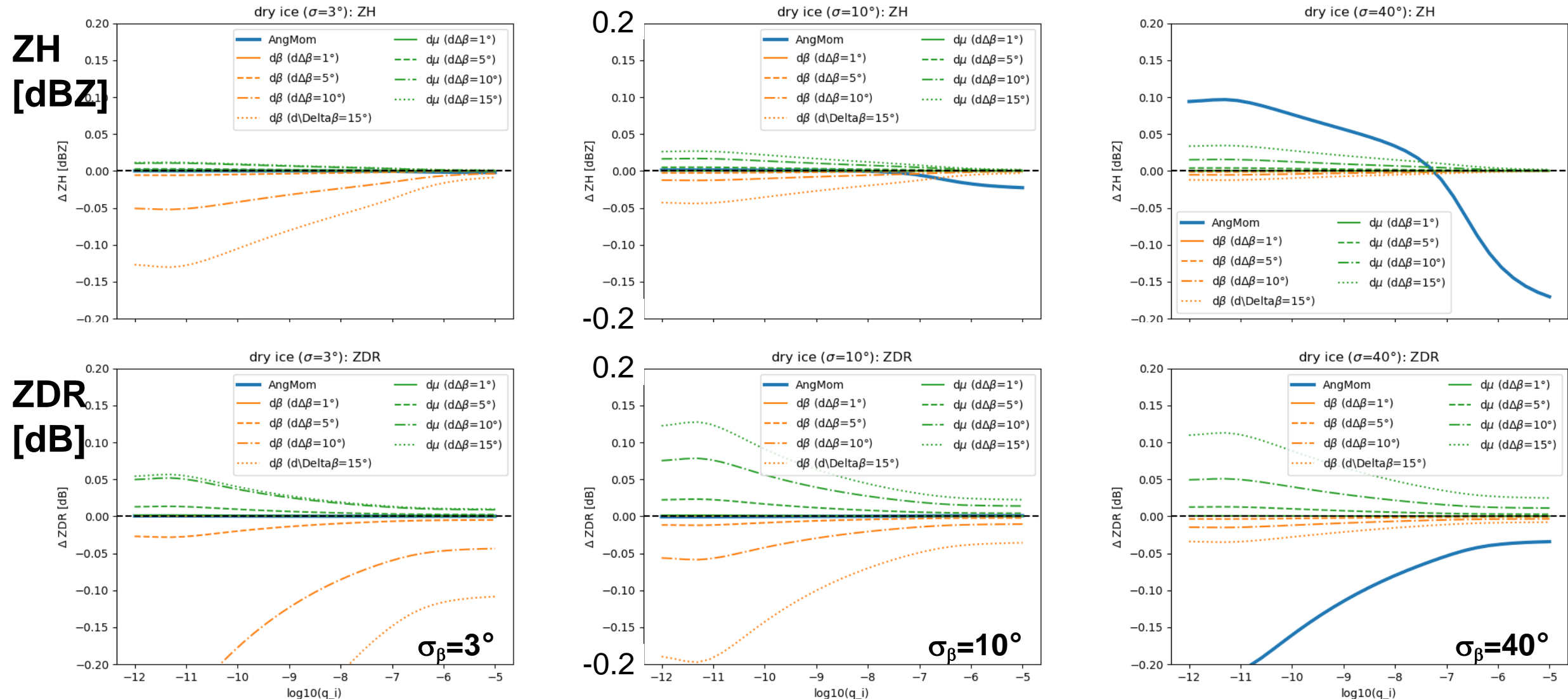
Consider any of them „more“ (physically) correct?



- From single particle tests:
- Simpson performs bad for polar angle integration
 - Trapezoidal with $\Delta\mu=\text{const}$ performs bad
 - Gauss-Legendre better than $\Delta\beta=\text{const}$ except for small σ_β

Explicit orientation averaging: Choice of dx & spacing

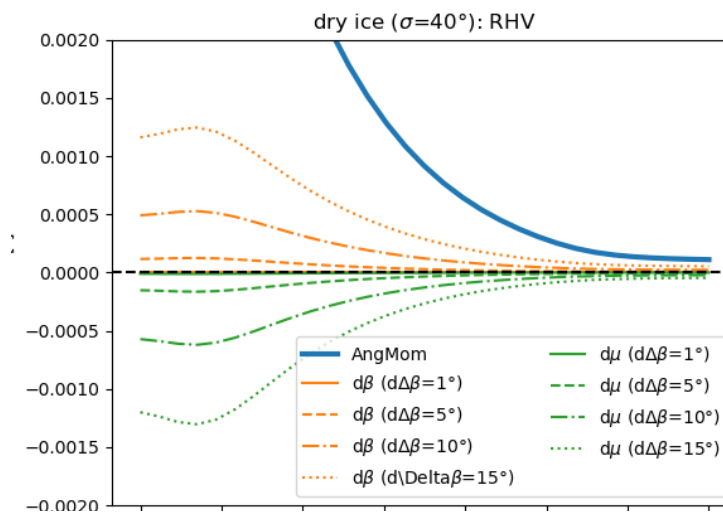
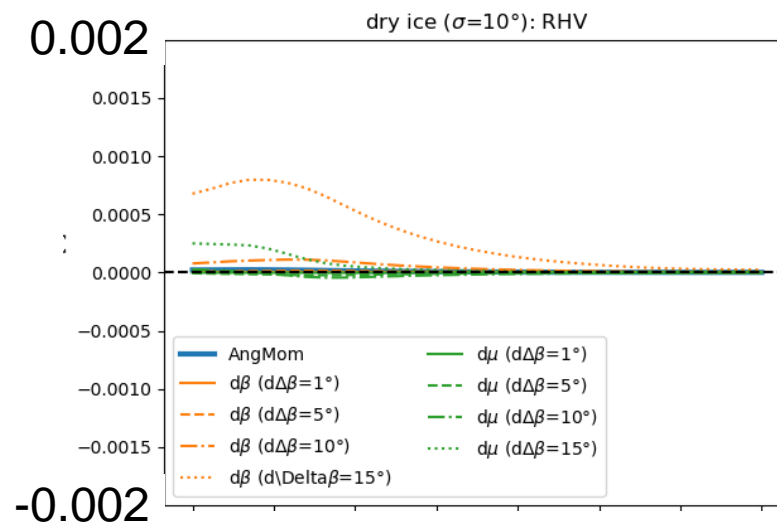
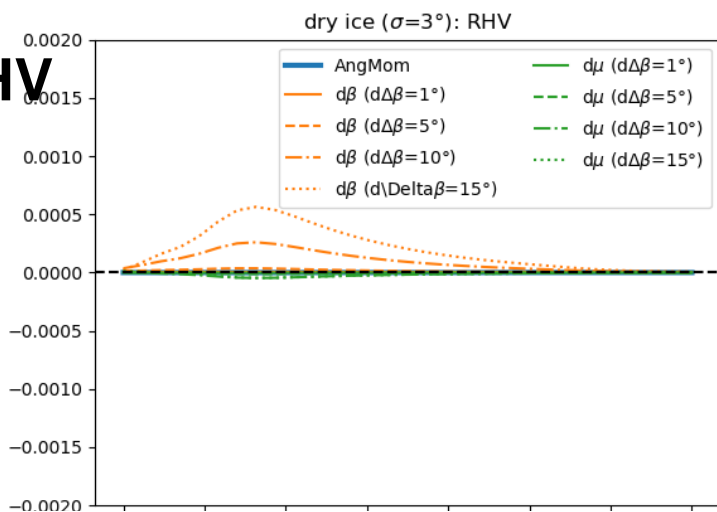
➔ Resulting EMVORADO bulk scattering LUT (dry ice example): Diff to **exp.orient.avg. high-res. reference**



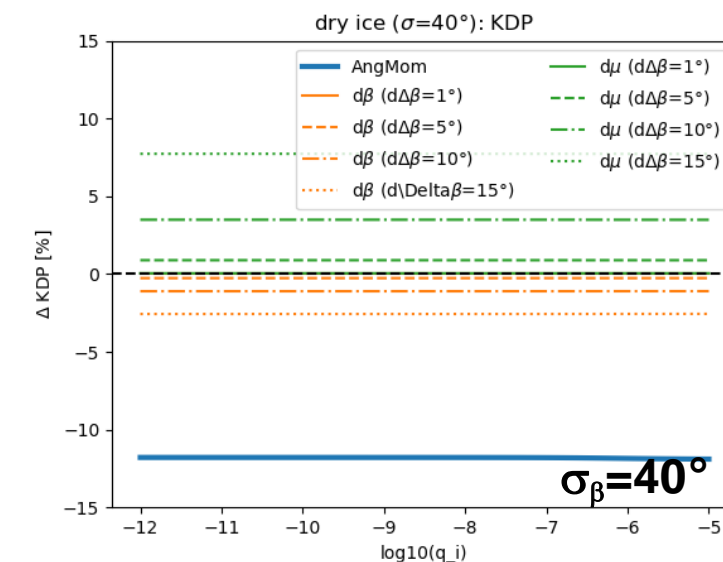
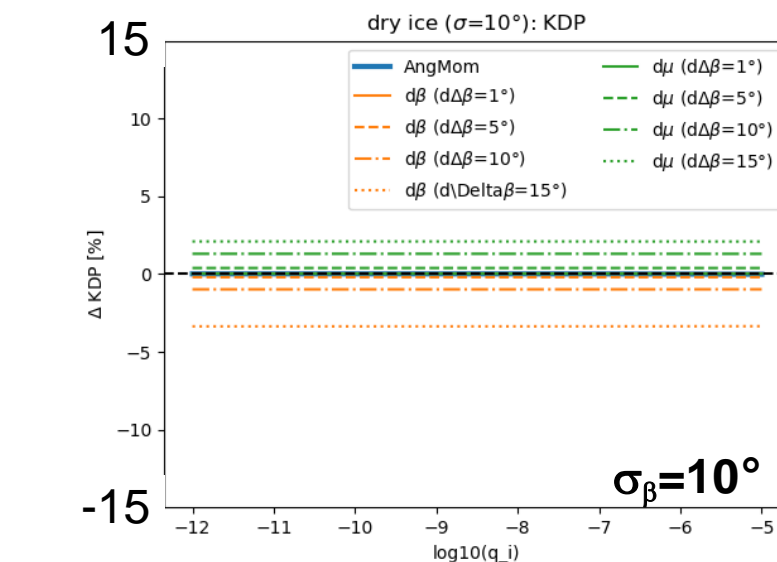
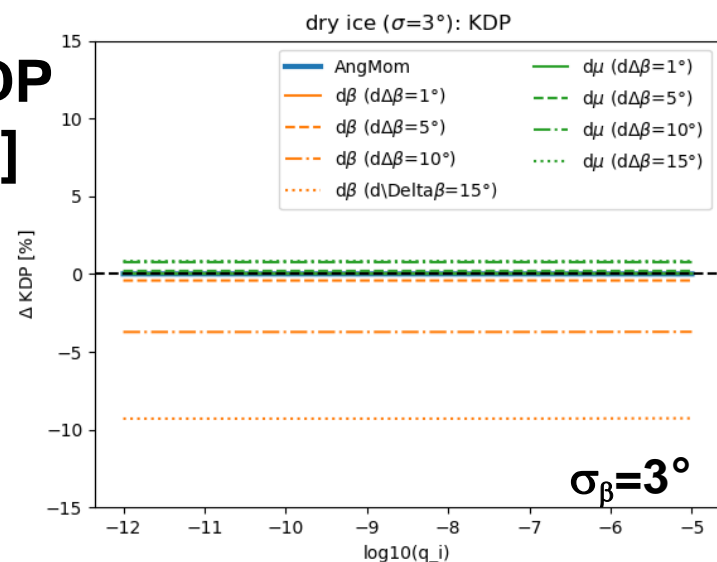
Explicit orientation averaging: Choice of dx & spacing

➔ Resulting EMVORADO bulk scattering LUT (dry ice example): Diff to **exp.orient.avg. high-res. reference**

RHV
[-]



KDP
[%]



→ From EMVORADO bulk LUT (dry ice example):

- $\int d\mu$ might mitigate $\beta=0^\circ$ non-contribution issue of $\int \sin(\beta) d\beta$
 - for $\mu(\Delta\beta=\text{const})$:
 - kinks at $\sigma_\beta < \Delta\beta$ removed in single particle example, but slightly larger offsets in general for same $\Delta\beta$
 - $d\mu$ clearly better at $\sigma_\beta < \Delta\beta$, similar performance at $\sigma_\beta \approx \Delta\beta$, slightly worse at $\sigma_\beta > \Delta\beta$
 - $d\mu$ overestimates, $d\beta$ underestimates
- for large(r) σ_β , any $\Delta\beta$ and dx better than **angmom** (clearly underestimates pol. moments)
 - **gridding over β on $\Delta\beta=\text{const}$ base that can easily be refined when/where needed**
 - $dx=d\mu$ or selection of dx dependent on σ_β**
 - partially refined grid or even mix of $d\mu$ (at small β) and $d\beta$ (at larger β) could be used**
- from dry ice LUT, $\Delta\alpha=30^\circ$ & $\Delta\beta=10^\circ$ **seem sufficient**
 - but should be further tested for rain & melting hydrometeors

(More) Questions?

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$\int z(\beta) p(\beta) \sin\beta \, d\beta \equiv \int z(\beta(\mu)) p(\beta(\mu)) \, d\mu$? **YES!**

$$\int_{x_1}^{x_2} f(x) \, dx = \int_{y_1(x_1)}^{y_2(x_2)} g(y) \, dy$$

$$f(x) \, dx = g(y) \, dy$$

$$g(y) = f(x) \frac{dx}{dy}$$

$$x = \cos\beta = \mu$$

$$\frac{dx}{d\cos\beta} = \frac{1}{\frac{d\cos\beta}{d\beta}} = \frac{1}{-\sin\beta}$$

$$\int g(y) \, dy = \int f(x) \frac{dx}{dy} \, dy = \int f(x) \frac{1}{-\sin\beta} \, d\cos\beta$$

$$= -f(x) \sin\beta \, d\cos\beta$$

$$f(x) = p(x) \cdot \sin\beta$$

$$g(y) = f(x) \frac{dx}{dy} = \frac{d\cos\beta}{d\beta} = -\sin\beta$$

$$= p(x) \cdot \sin\beta \cdot (-\sin\beta)$$

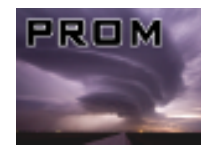
$$= -p(x)$$

$$\int_0^1 p(x) \, d\cos\beta$$

$$y_1 = \cos\beta_1 = \cos 0^\circ = 1$$

$$y_2 = \cos\beta_2 = \cos 90^\circ = 0$$

$$= \int_1^0 p(x) \, d\cos\beta$$

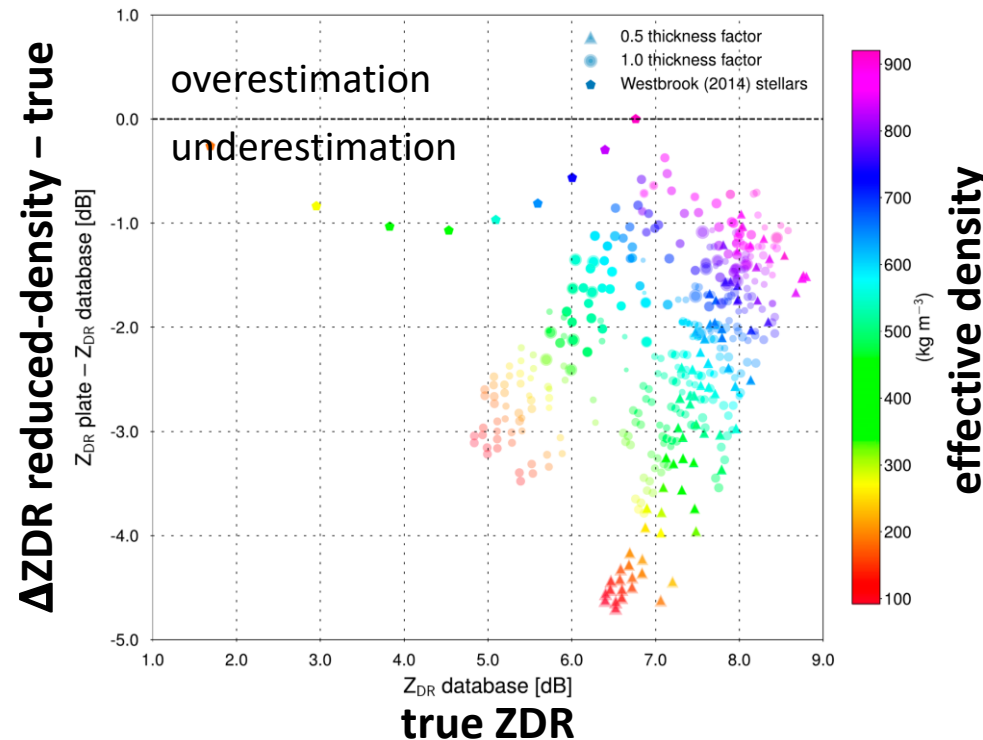


Probably the most popular approach to setup particles consistent to model constraints (keeping m , D , and aspect ratio unchanged) with T-Matrix suitable shapes.

T-Matrix based simulations show a **consistent deficit** in terms of **polarimetric response** in the dendritic growth layer where large, “fluffy” particles prevail.

Schrom & Kumjian (2018)

- assessed errors in polarimetric scattering properties of homogeneous reduced-density particles as proxies of branched planar crystals (both from DDA)
- found persistent underestimation of ZDR, the worse the less dense
- provided detailed explanation for the **role of internal structure** from dipole interactions

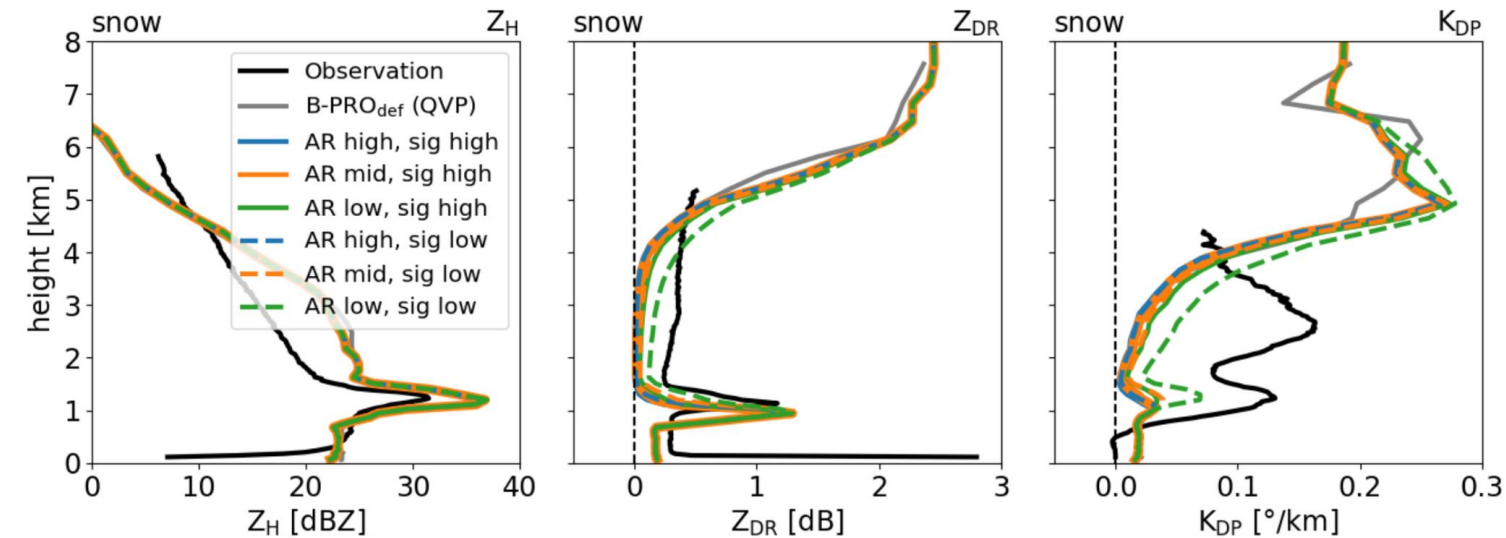


... consistent deficit in terms of polarimetric response ...

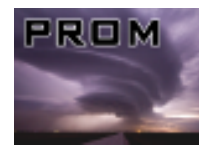
There are further explanations for lack of polarimetric signals!

FO uncertainties that can contribute include, e.g.,

- melting models
- dielectric properties (primarily of air-ice(-water) mixtures)
- *shape and orientation assumptions*



snow	
AR _{high}	Ryzhkov et al. (2011) $\max(1.0 - 20D, 0.8)$
AR _{mid}	Xie et al. (2016) $\max(0.7 - 10D, 0.5)$
AR _{low}	Dunnavan et al. (2019) 0.4
σ_{high}	Ryzhkov et al. (2011) 40°
σ_{low}	Matsui et al. (2019) 20°



... consistent deficit in terms of polarimetric response ...

There are **further explanations & reasons** for lack of polarimetric signals!

FO uncertainties that can contribute include, e.g.,

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- shape and orientation assumptions

Regarding **model microphysics** these include, e.g.,

- hydrometeor size distribution
- hydrometeor class partitioning
 - o lack of secondary ice
 - o wet growth processes
- mass-size relation
- mixed-phase hydrometeors

→ **Can we draw robust conclusions about model microphysics from synthetic signals based on homogeneous particle approaches?**