

Rheinische -riedrich-Wilhelms-Universität Bonn



### Climate model PArameterizations informed by RAdar (PARA)

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#### **Overall motivation of PARA**

The overall goal of PARA is to evaluate and improve the representation of microphysical processes in ICON-GCM exploiting high-resolved polarimetric radar observations.



ICON GCM ~ 150km

Radars  $\sim$  0.15 – 1 Km  $_2$ 

#### Work Programme

- The project focuses on the following processes:
- Spatial sub-grid scale variability of cloud ice
   Aggregation
   Melting of precipitation
   Evaporation

**<u>Note</u>**: The smallest size particle which is assigned to snow in ICON model is 100µm.

### Spatial sub-grid scale variability of cloud ice (Uni Bonn)

- Our goal is to provide the mean value and variance of:
- Ice Water Content IWC (Ice + Snow + Overall)
- **\***Total number Concentration N<sub>t</sub> (Ice + Snow + Overall)
- **\therefore** Mean Volume Diameter  $D_m$  (Ice + Snow + Overall)

as a function of height, using polarimetric radar measurements and applying the most recent ice microphysical retrievals by Ryzhkov et al. (2018) and Bukovcic et al. (2020).

#### Data and Methodology

#### Data



Polarimetric radar data from the:

- ➢ Radar network of DWD with C − band weather radars.
- ➢ BoXPol X − band radar
- ►JuXPol X band radar



#### Methodology - Ice microphysical retrievals

This study uses the most recent ice microphysical retrievals by Ryzhkov et al. (2018) and Bukovcic et al. (2020).

1) 
$$D_m(mm) = = -0.1 + 2.0(\frac{Z_{DP}}{K_{DP}\lambda})^{1/2}$$
  
2)  $IWC(\frac{gr}{m^3}) = 4 \times 10^{-3} \frac{K_{DP}\lambda}{1-Zdr^{-1}}$   
3)  $\log(N_t)(\frac{1}{L}) = 0.1Z - 2\log(0.2D_m^2) - 1.33$   
Ryzhkov et al. (2018)  
 $D_m(mm) = 0.7(\frac{Z}{K_{DP}\lambda})^{1/3}$   
 $IWC(\frac{gr}{m^3}) = 0.3(K_{DP}\lambda)^{0.66}Z^{0.28}$   
 $\log(N_t)(\frac{1}{L}) = 0.1Z - 2\log(0.2D_m^2) - 1.33$ 

*Note*: σ= 0 and a/b=0.65

#### Methodology – Statistical Errors

Estimation of the statistical errors of  $Z_{DR}$ ,  $Z_H$  and  $K_{DP}$  following **Ryzhkov and Zrnic (2018)** and **Vulpiani et al. (2012).** 



Using the Gaussian Error propagation is estimated the resulting error in  $IWC(K_{DP}, Z)$  and  $IWC(Z_{DR}, K_{DP})$ for each radar bin.

- ➢Averaging is needed in order to be reduced the statistical error of the retrievals.
- This study uses the Quasi Vertical Profiles(QVP) methodology and the QVP methodology applied in azimuthal sectors.

### Methodology – Quasi Vertical Profiles (QVPs)



- Low statistical errors.
- Coarse resolution which cannot provide alone the sub-grid scale ice variability.

Fig. 2) Quasi Vertical Profile methodology

## Methodology – QVPs applied in azimuthal sectors

- + Provides the horizontal variability in azimuths.
- Higher statistical errors.

Challenge:

How?

The appropriate choice of the sector size.

Comparisons between the variability of IWC (of both formulas) in azimuths with the mean error of IWC for different sector sizes

#### Stdev(IWC) vs Error(IWC)

Fig. 3) Variability and error with 10 increasing sector size.



# Methodology - Ice and snow component of IWC, $D_m$ , $N_t$



Fig. 4) QVPs of Total Concentration for the rain event 12/04/2013 using BoXPol radar and elevation angle 18 degrees.



### Results (snow component)



Fig. 5) Height profiles of IWC(snow), Nt(snow) and Dm(snow) for the rain events 12/04/2013, 07/10/2014 and 16/11/2014 using BoXPol radar and elevation angle 18 degrees.

### Results (ice component)



Fig. 6) Height profiles of IWC(ice), Nt(ice) and Dm(ice) for the rain events 12/04/2013, 07/10/2014 and 16/11/2014 using BoXPol radar and elevation angle 18 degrees.

#### Summary



#### Future work

More cases.

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Study of the snow formation with aggregation